

Reflections on Geomorphometry and Hydrology

David Tarboton

Utah State University

<http://hydrology.usu.edu/taudem>



Utah Water Research Laboratory
UtahStateUniversity

<http://hydrology.usu.edu>



David
Maidment



Teklu
Tesfa



Irene
Garousi-Nejad



Shaowen
Wang



Matt
Baker



Xing
Zheng



Bob
Pack



Kiran
Chinnayakanahalli



Paolo
Tarolli



Rob
Wallace



Yan
Liu



Dan
Watson



Kim
Schreuders



Ahmet
Yildirim



Dan
Ames



Pabitra
Dash



Ajay
Prasad

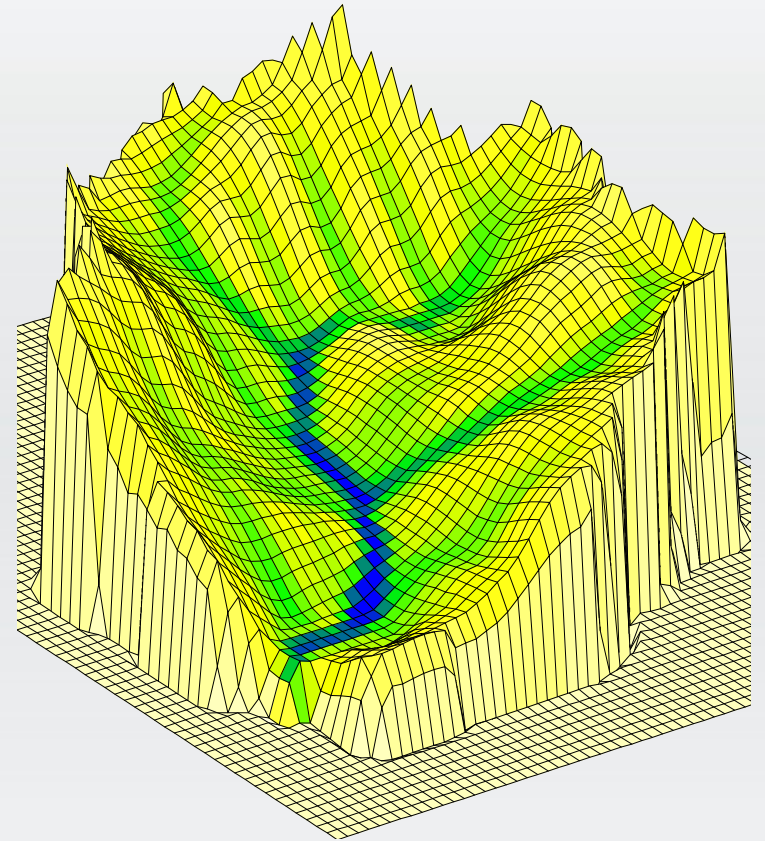


Utah Water Research Laboratory
Utah State University

<http://hydrology.usu.edu>

Outline

- Watershed and channel network delineation and objective channelization threshold determination
- D-infinity and multiple flow directions
- Generalized terrain-based flow analysis (flow algebra)
- Height above nearest drainage (HAND) and flood inundation mapping
- TauDEM software and parallel programming (speeding it up)



Topography and Hydrology

- Watersheds are fundamentally the most basic hydrologic landscape elements
- Topography dictates the flow of water across the landscape
- Flowing water sculpts the landscape
- This synergy is at the heart of much hydrologic modeling relating to questions of runoff generation important for flooding and water resources
- Representing hydrologic processes at high resolution is important to help solve these problems

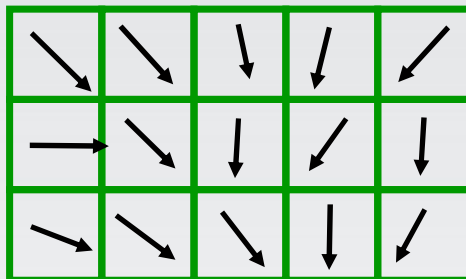


General Terrain Flow Model

Input Digital Elevation Model



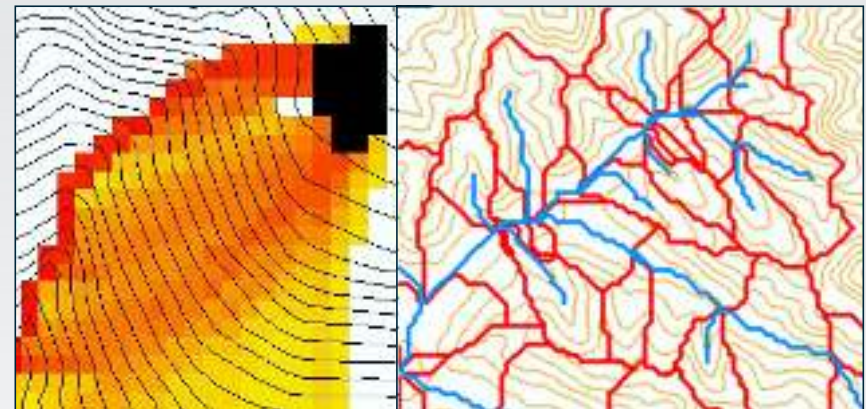
Flow Field



Hydrologic Conditioning (Often Filling)



Flow Related Terrain Information



“In the most simple terms, geomorphometry is everything you can extract from DEMs”

www.geomorphometry.org

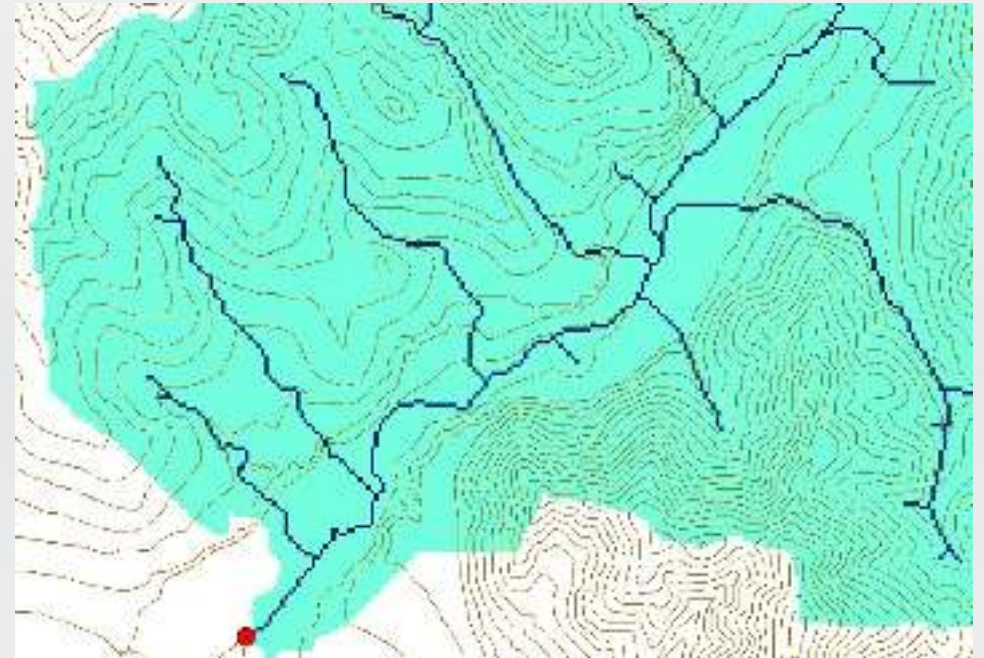


Utah Water Research Laboratory
Utah State University

<http://hydrology.usu.edu>

Basic Watershed and Channel Network Delineation

- Pit Removal (Filling)
- Flow direction calculation (D8)
- Flow accumulation (may be weighted)
- Channel (Stream) definition using a threshold
- Channels and watersheds (catchments/basins)



Conditioning the DEM - Pit Filling

Increase elevation until the pit drains to a neighbor and ultimately out of the domain



Original
DEM

7	7	6	7	7	7	7	5	7	7
9	9	8	9	9	9	9	7	9	9
11	11	10	11	11	11	11	9	11	11
12	12	8	12	12	12	12	10	12	12
13	12	7	12	13	13	13	11	13	13
14	7	6	11	14	14	14	12	14	14
15	7	7	8	9	15	15	13	15	15
15	8	8	8	7	16	16	14	16	16
15	11	11	11	11	17	17	6	17	17
15	15	15	15	15	18	18	15	18	18

Pits

Grid cells or zones completely
surrounded by higher terrain

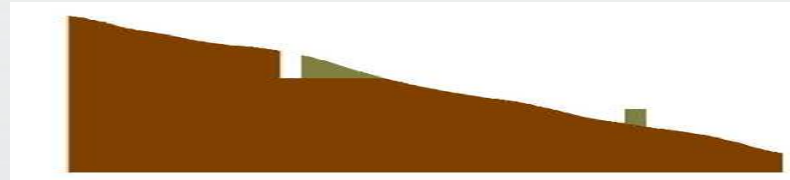
Pits Filled

7	7	6	7	7	7	7	5	7	7
9	9	8	9	9	9	9	7	9	9
11	11	10	11	11	11	11	9	11	11
12	12	10	12	12	12	12	10	12	12
13	12	10	12	13	13	13	11	13	13
14	10	10	11	14	14	14	12	14	14
15	10	10	10	10	15	15	13	15	15
15	10	10	10	10	16	16	14	16	16
15	11	11	11	11	17	17	14	17	17
15	15	15	15	15	18	18	15	18	18



Conditioning the DEM - Carving

Lower elevation of neighbor along a predefined drainage path until the pit drains to the outlet point



Original
DEM

7	7	6	7	7	7	7	5	7	7
9	9	8	9	9	9	9	7	9	9
11	11	10	11	11	11	11	9	11	11
12	12	8	12	12	12	12	10	12	12
13	12	7	12	13	13	13	11	13	13
14	7	6	11	14	14	14	12	14	14
15	7	7	8	9	15	15	13	15	15
15	8	8	8	7	16	16	14	16	16
15	11	11	11	11	17	17	6	17	17
15	15	15	15	15	18	18	15	18	18

Pits

Carved
DEM

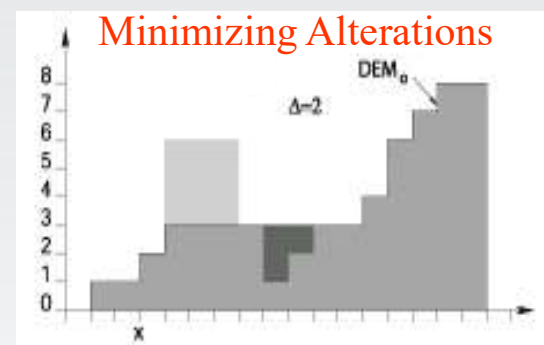
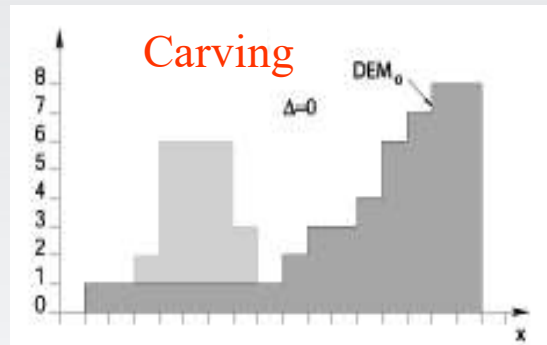
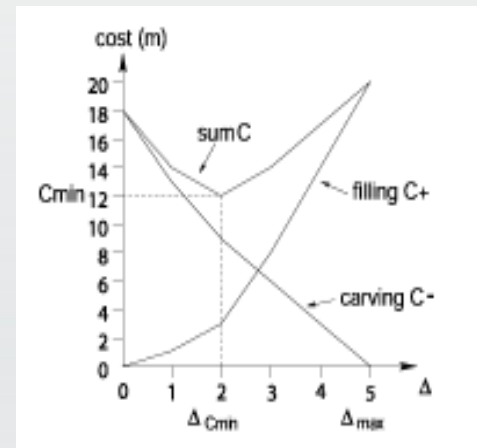
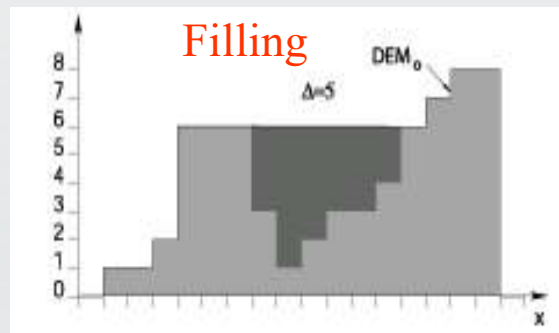
7	7	6	7	7	7	7	5	7	7
9	9	6	9	9	9	9	6	9	9
11	11	6	11	11	11	11	6	11	11
12	12	6	12	12	12	12	6	12	12
13	12	6	12	13	13	13	6	13	13
14	7	6	11	14	14	14	6	14	14
15	7	7	7	9	15	15	6	15	15
15	8	8	8	7	16	16	6	16	16
15	11	11	11	11	17	17	6	17	17
15	15	15	15	15	18	18	15	18	18

Carve outlets



Optimal removal of spurious pits in grid digital elevation models

Pierre Soille, WRR 2004, <http://doi.org/10.1029/2004WR003060>



Conditioning the DEM – Optimally adjusted

Original
DEM

7	7	6	7	7	7	7	5	7	7
9	9	8	9	9	9	9	7	9	9
11	11	10	11	11	11	11	9	11	11
12	12	8	12	12	12	12	10	12	12
13	12	7	12	13	13	13	11	13	13
14	7	6	11	14	14	14	12	14	14
15	7	7	8	9	15	15	13	15	15
15	8	8	8	7	16	16	14	16	16
15	11	11	11	11	17	17	6	17	17
15	15	15	15	15	18	18	15	18	18

Pits

Carved

7	7	6	7	7	7	7	5	7	7
9	9	7	9	9	9	9	7	9	9
11	11	7	11	11	11	11	9	11	11
12	12	7	12	12	12	12	10	12	12
13	12	7	12	13	13	13	11	13	13
14	7	7	11	14	14	14	12	14	14
15	7	7	7	9	15	15	13	15	15
15	8	8	8	7	16	16	14	16	16
15	11	11	11	11	17	17	14	17	17
15	15	15	15	15	18	18	15	18	18

Optimally
adjusted

Filled

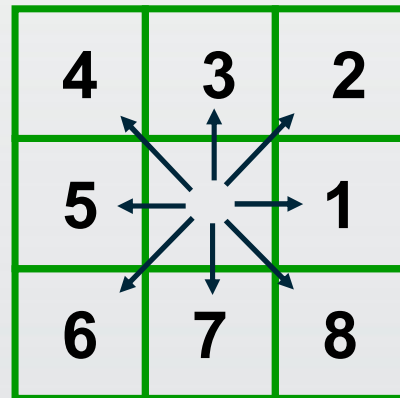


D8 Flow Direction Model

Direction of steepest descent

30

80	74	63
69	67	56
60	52	48

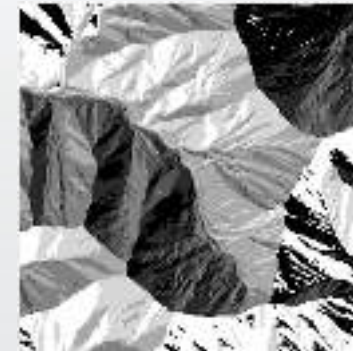


$$\frac{67 - 52}{30} = 0.50$$

$$\frac{67 - 48}{30\sqrt{2}} = 0.45$$

Slope = Drop/Distance
Steepest down slope direction

Note that this is different from the 3x3 or 5x5 slope that Peter Guth discussed Monday. It only looks down to where water may flow.



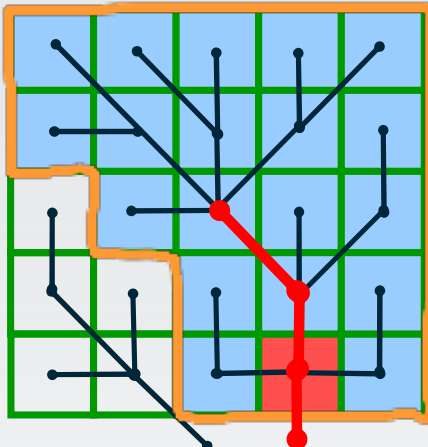
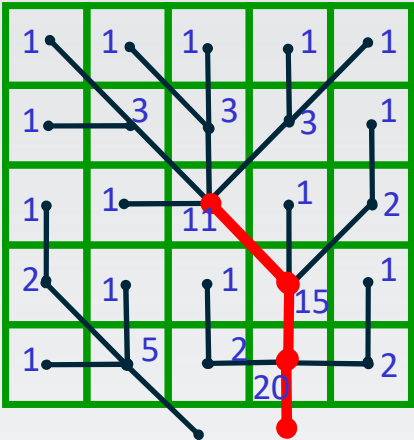
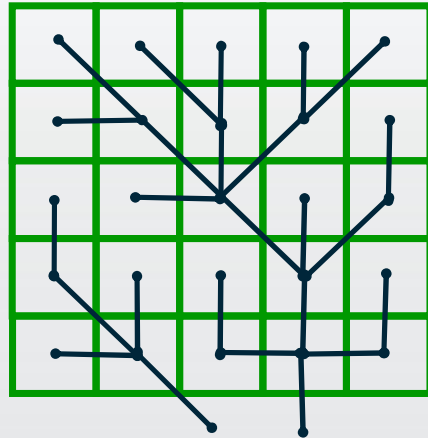
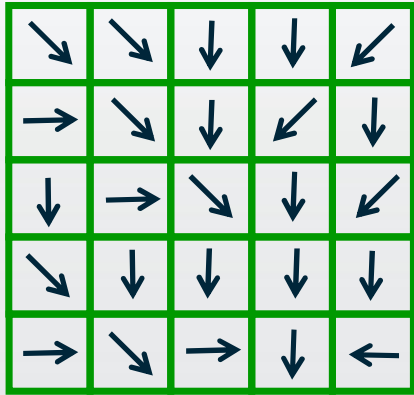
O'Callaghan, J. F. and D. M. Mark, (1984), [https://doi.org/10.1016/S0734-189X\(84\)80011-0](https://doi.org/10.1016/S0734-189X(84)80011-0).



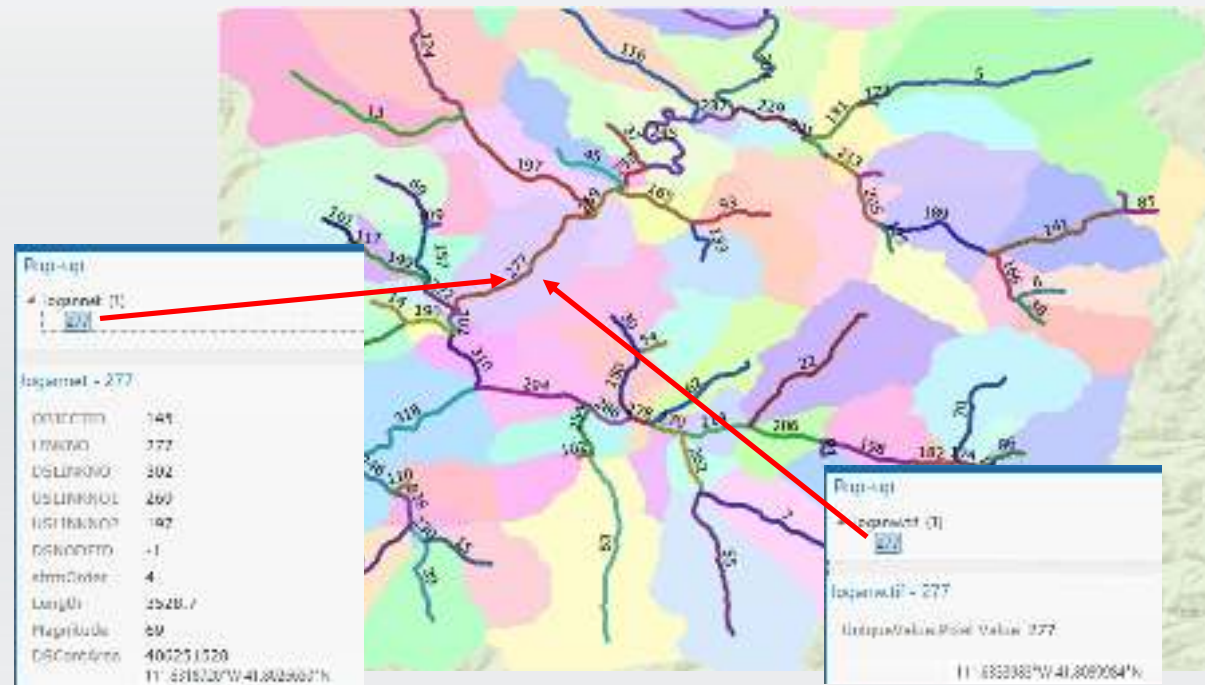
Utah Water Research Laboratory
Utah State University

<http://hydrology.usu.edu>

Grid Network, Flow Accumulation, Streams and Watersheds



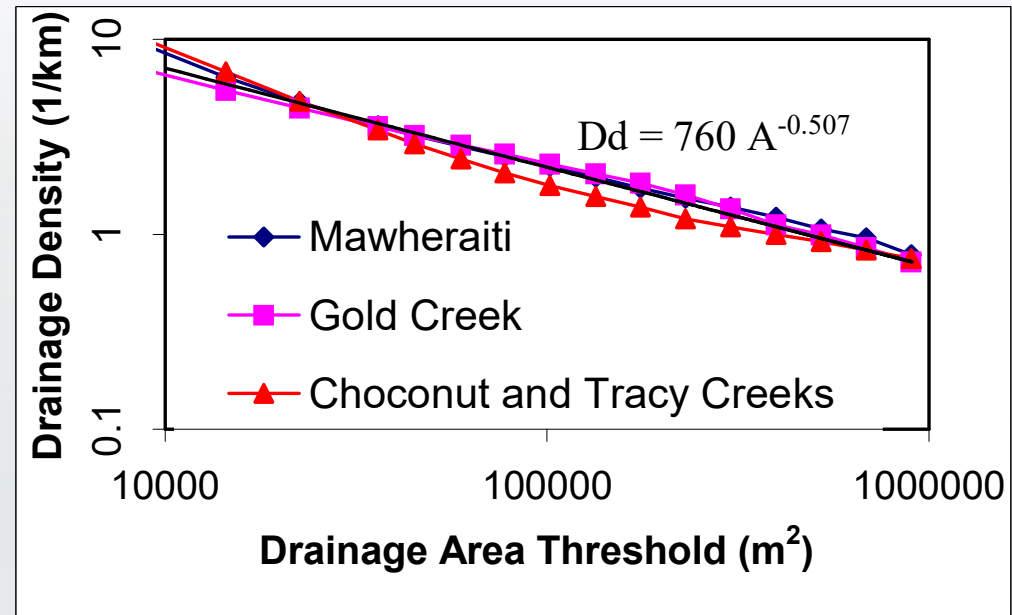
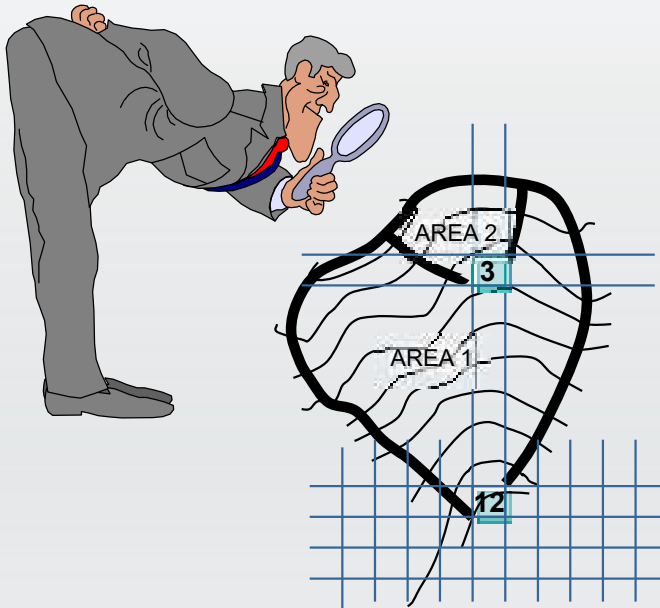
Streams linked to associated catchments



The starting point for catchment based distributed hydrologic modeling



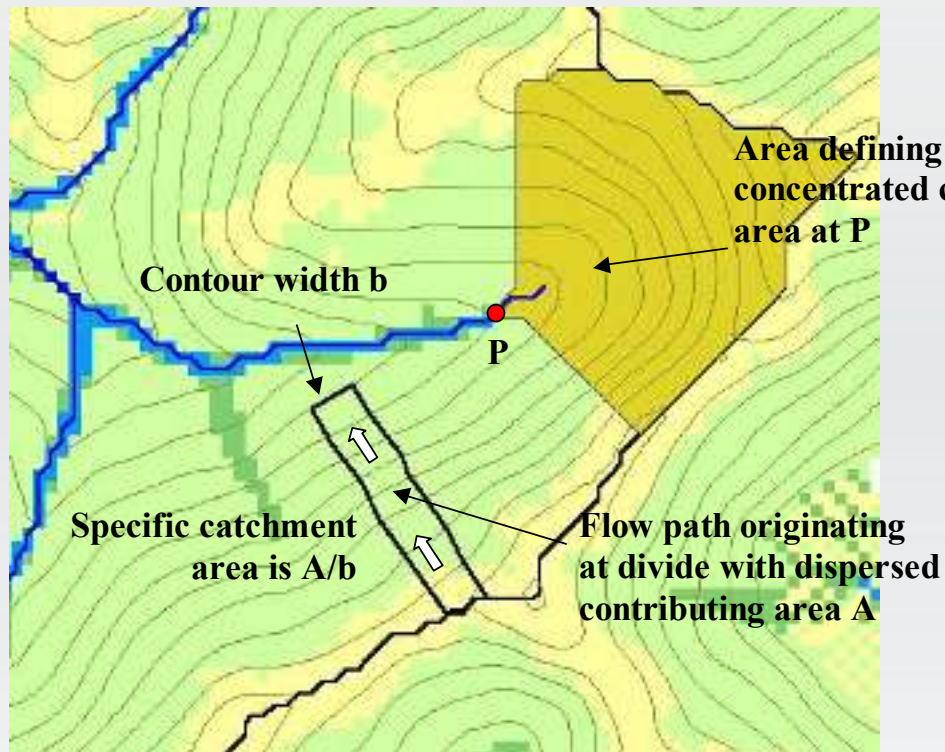
Where do Channels Begin?



Drainage density (total channel length divided by drainage area) as a function of drainage area support threshold used to define channels for the three study watersheds.

How to decide on stream delineation threshold ?
Why is this important ?

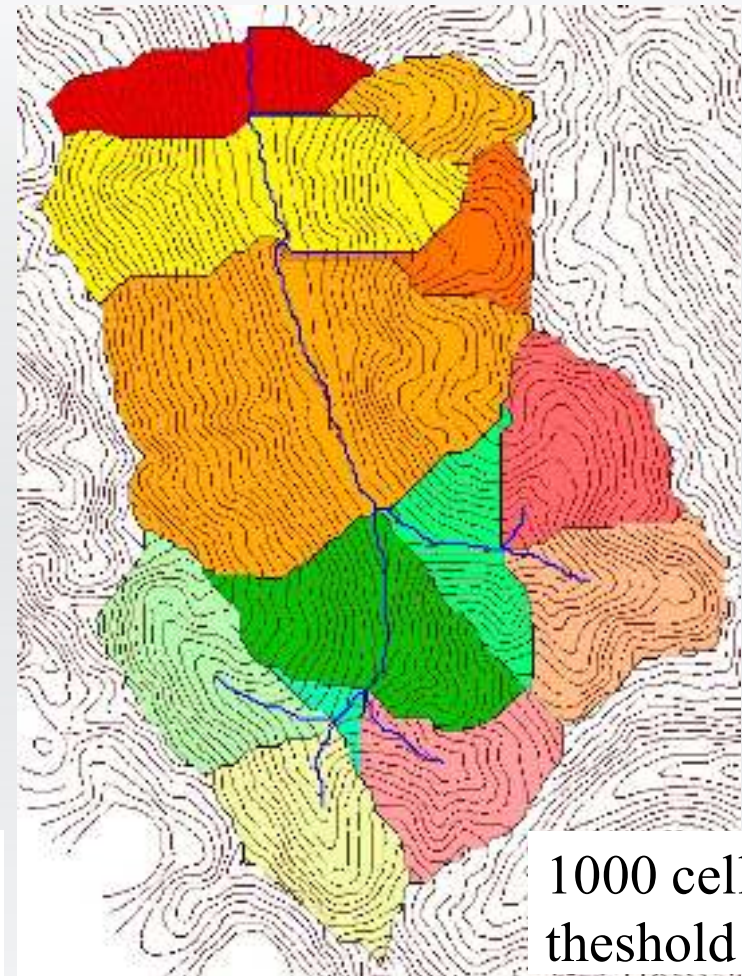
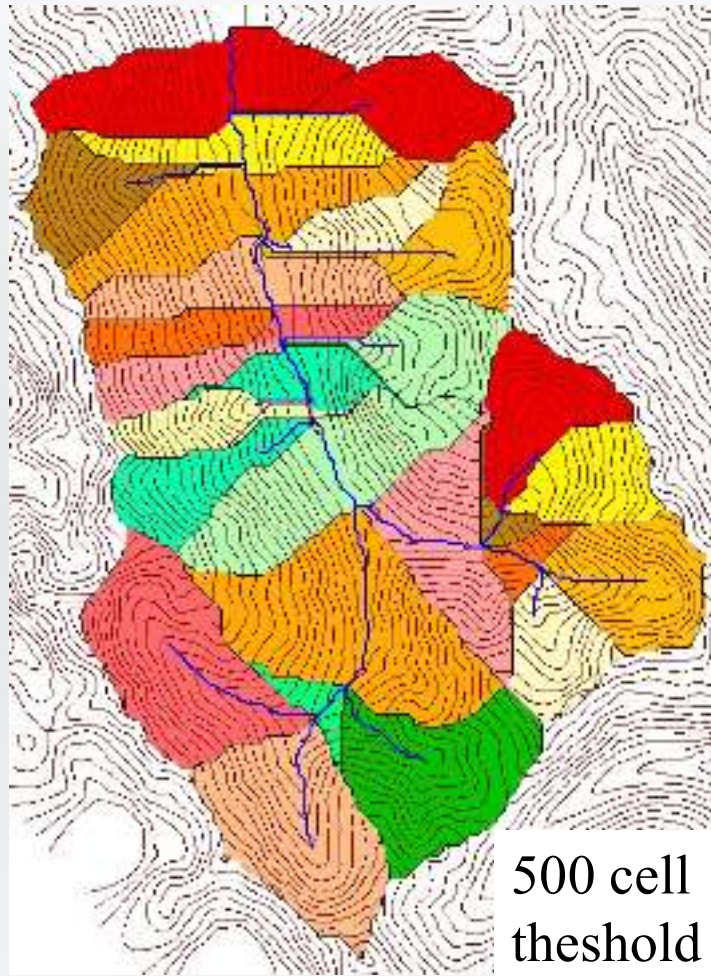
Hydrologic processes are different on hillslopes and in channels. It is important to recognize this and account for this in the delineation of streams.



Drainage area can be concentrated or dispersed (specific catchment area) representing concentrated or dispersed flow.



The scale of catchment model elements is sensitive to channelization support area threshold



Alternative views

Although the river and hill-side waste do not resemble each other at first sight, they are only the extreme members of a continuous series, and when this generalization is appreciated, one may fairly extend the “river” all over its basin and up to its very divides. Ordinarily treated, the river is like the veins of a leaf; broadly viewed it is like the entire leaf.

Davis, W. M., (1899), "The geographical cycle," Geogr. J., 14: 481-504
(reproduced in Geographical Essays, edited by W. M. Davis, Ginn, Boston, 1909).

landscape dissection into distinct valleys is limited by a threshold of channelization that sets a finite scale to the landscape.

Montgomery, D. R. and W. E. Dietrich, (1992), "Channel Initiation and the Problem of Landscape Scale," Science, 255: 826-830.

Examples of differently textured topography



Badlands in Death Valley.
from Easterbrook, 1993, p 140.



Coos Bay, Oregon Coast Range.
from W. E. Dietrich



Gently Sloping Convex Landscape



From W. E. Dietrich



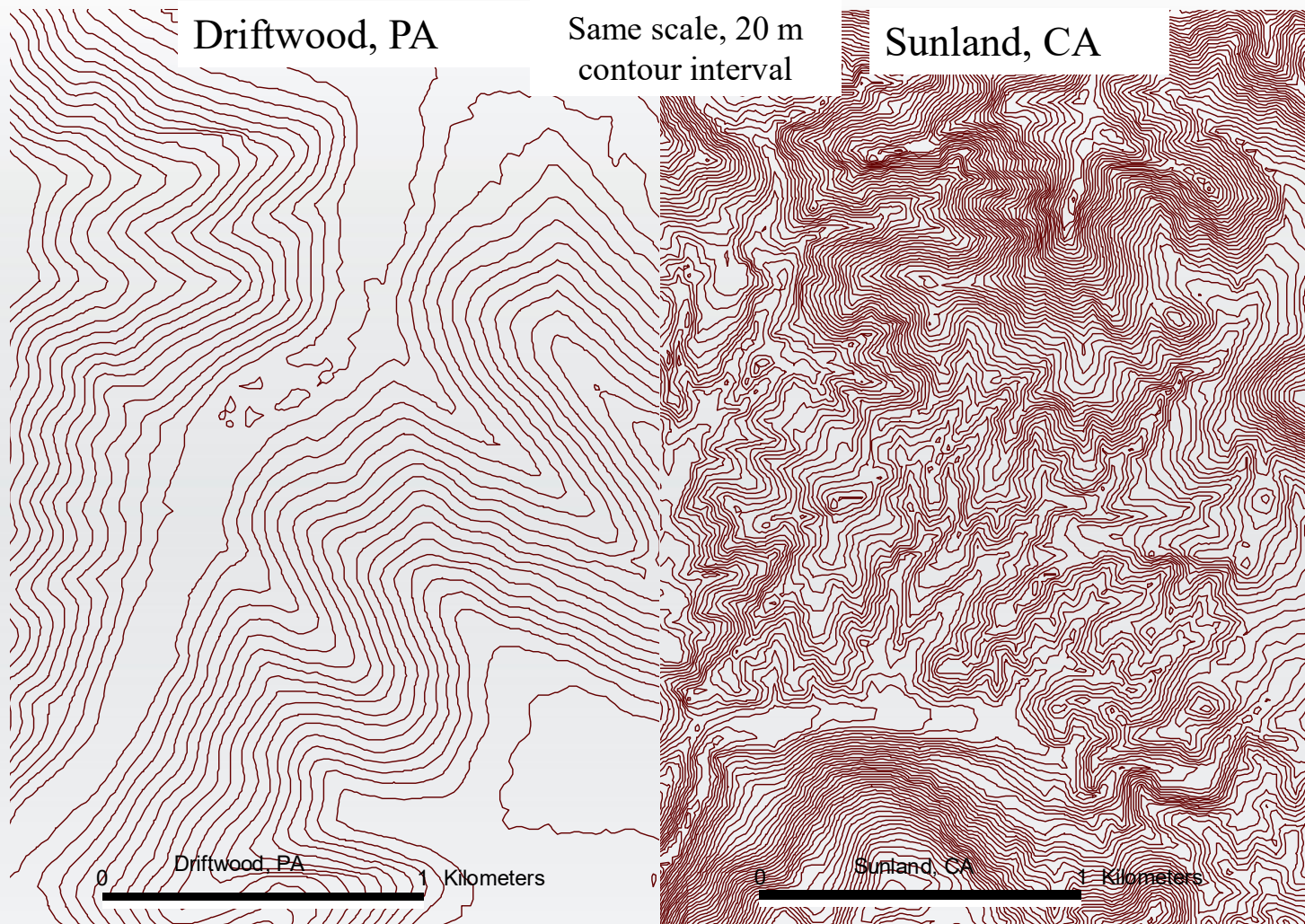
Utah Water Research Laboratory
Utah State University

<http://hydrology.usu.edu>



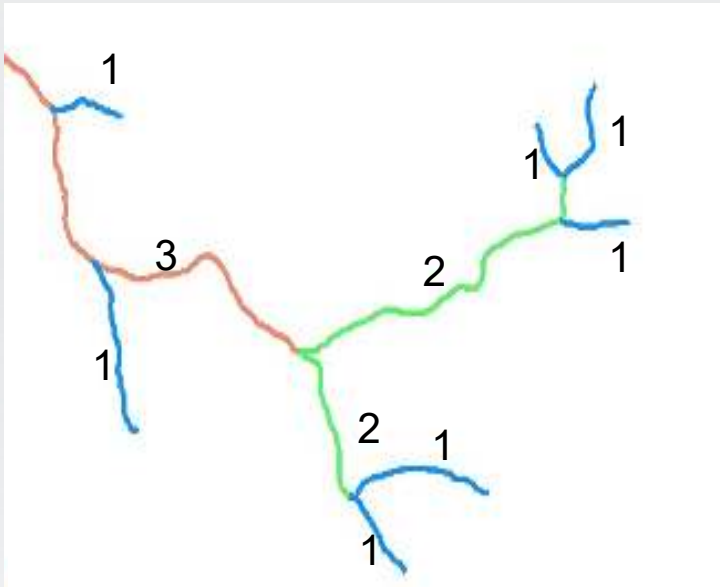


Contours indicate topographic texture

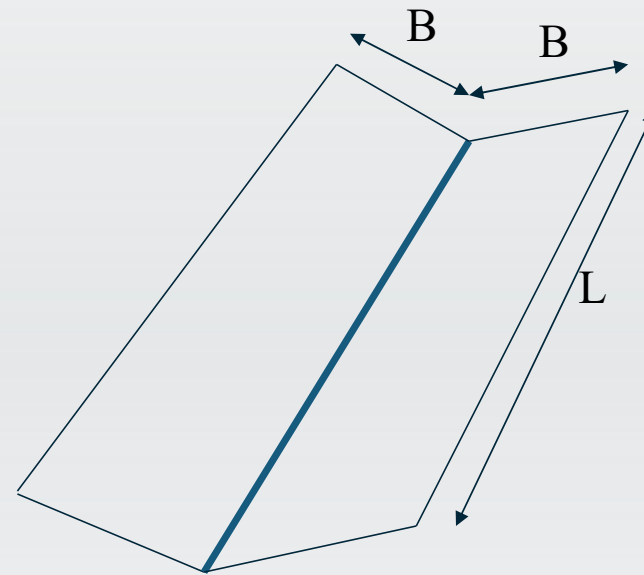


Topographic texture, drainage density and “laws” of geomorphology

Horton/Strahler stream ordering
Stream networks obey (approximately) the empirically established Hortons “laws”



Drainage density “measures” the scale of topographic texture



$$D_d = L/A$$

$$\text{Hillslope length} = B$$

$$A = 2B L$$

$$D_d = L/A = 1/(2B)$$

$$\therefore B = 1/(2D_d)$$

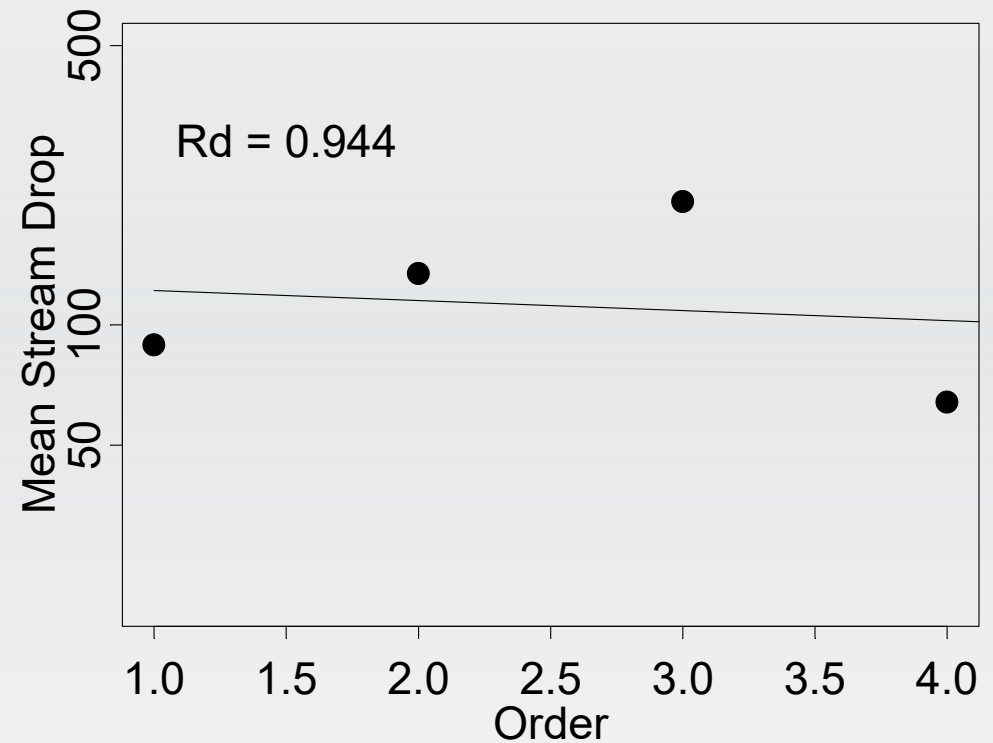
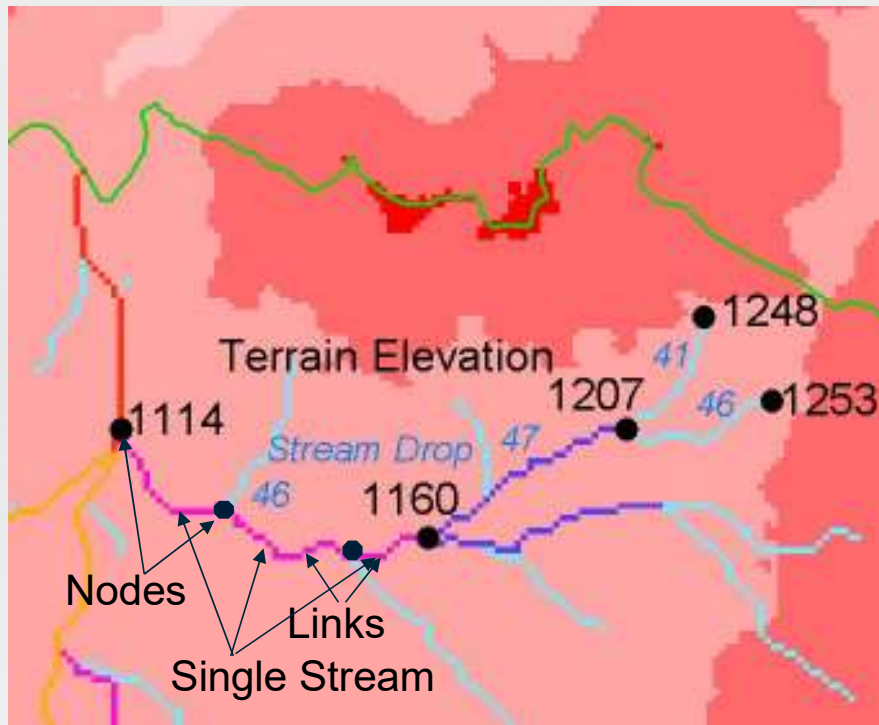
$$\text{Hillslope length} \approx 1/(2D_d)$$

Related concepts: Topographic grain, contour crenulations, ruggedness, topographic roughness



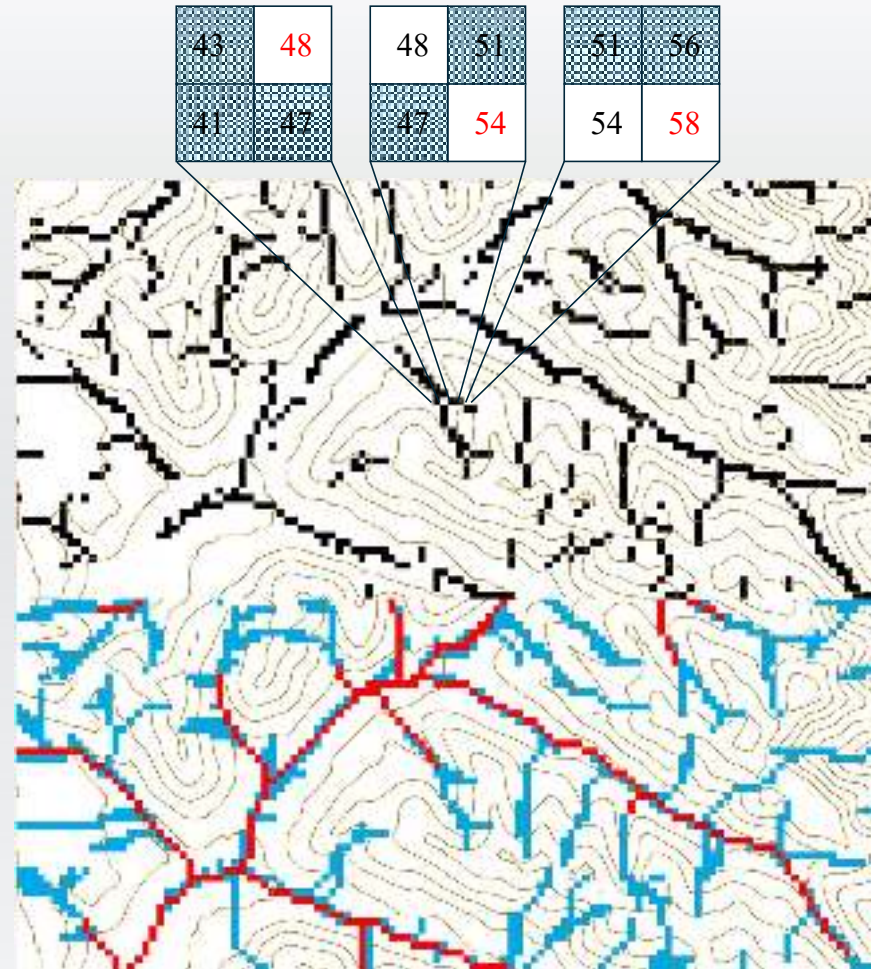
Constant Stream Drops Law

Broscoe, A. J., (1959), "Quantitative analysis of longitudinal stream profiles of small watersheds," Office of Naval Research, Project NR 389-042, Technical Report No. 18, Department of Geology, Columbia University, New York.



Approaching the objective mapping of channel networks

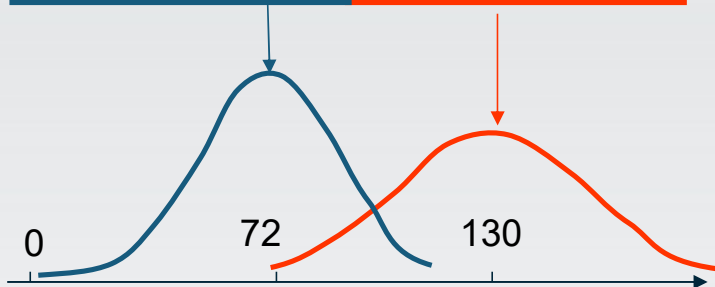
- Adjust for spatially variable drainage density
- Local valley grid cells (Peucker and Douglas, 1975, Comput. Graphics Image Proc. 4:375)
- Connect by taking valley grid cells as weights for flow accumulation
- Threshold the weighted flow accumulation grid



Select threshold to map channel networks at the finest resolution consistent with empirical constant drop “law”

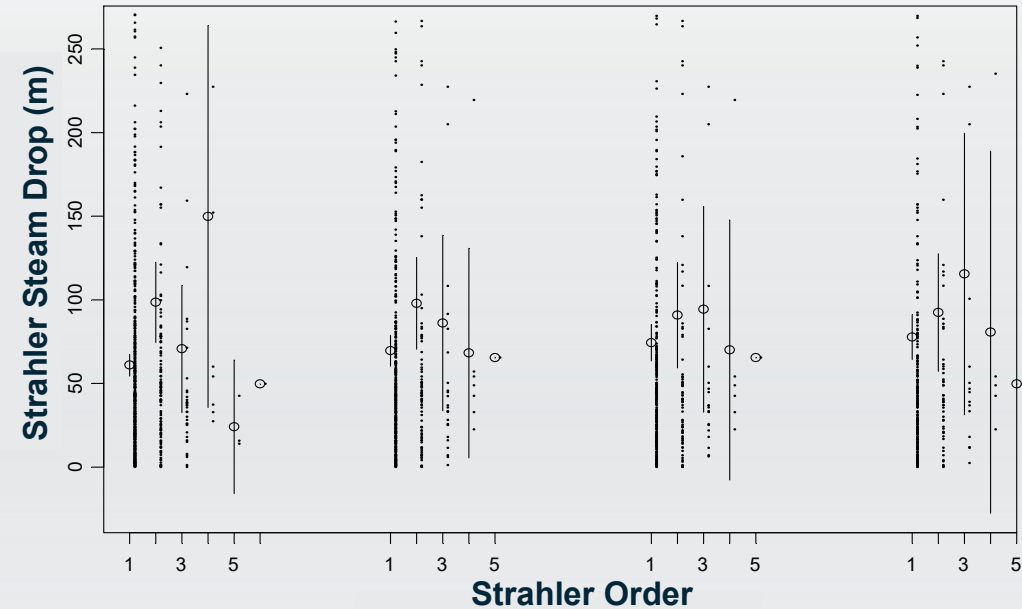
t-test for difference in mean between first and higher order stream drops

	Order 1		Order 2-4
Mean X	72.2	Mean Y	130.3
Std X	68.8	Std Y	120.8
Var X	4740.0	Var Y	14594.5
Nx	268	Ny	81



Use smallest threshold for which $|t| < 2$ ($p=0.05$) to map channels

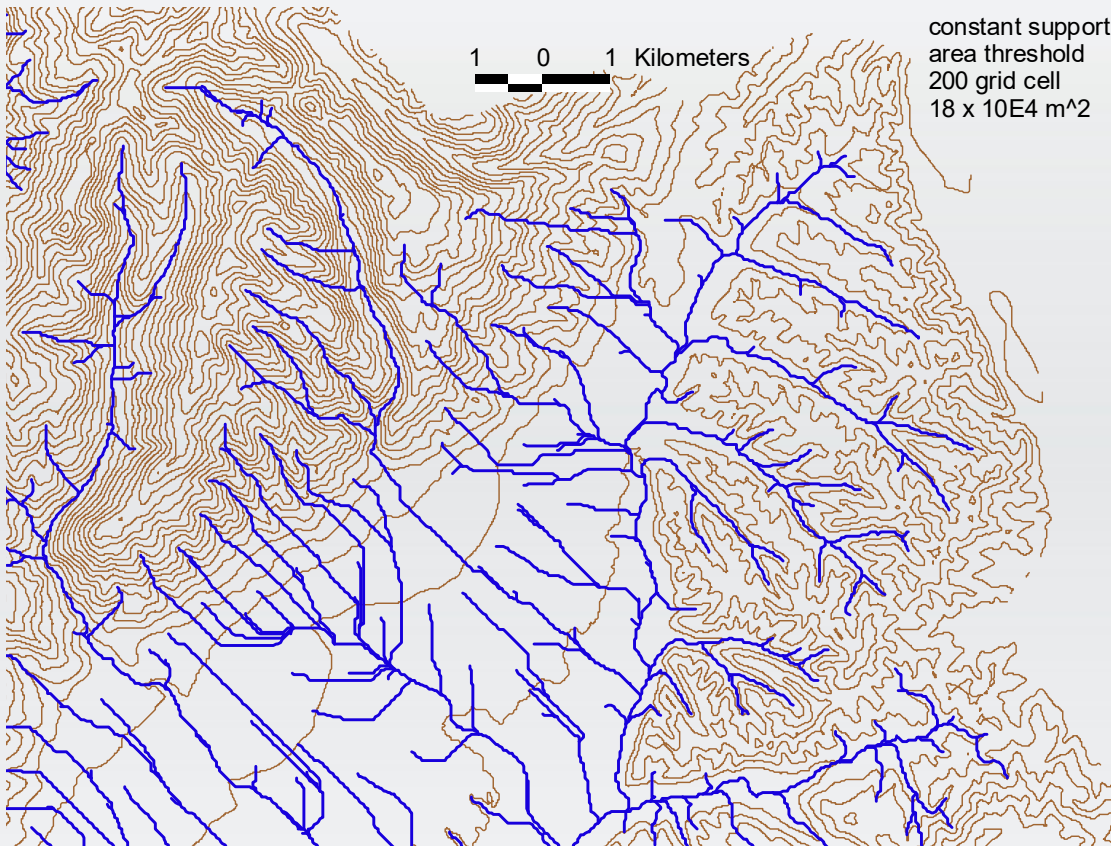
Weighted Valley Cell Contributing Area Threshold



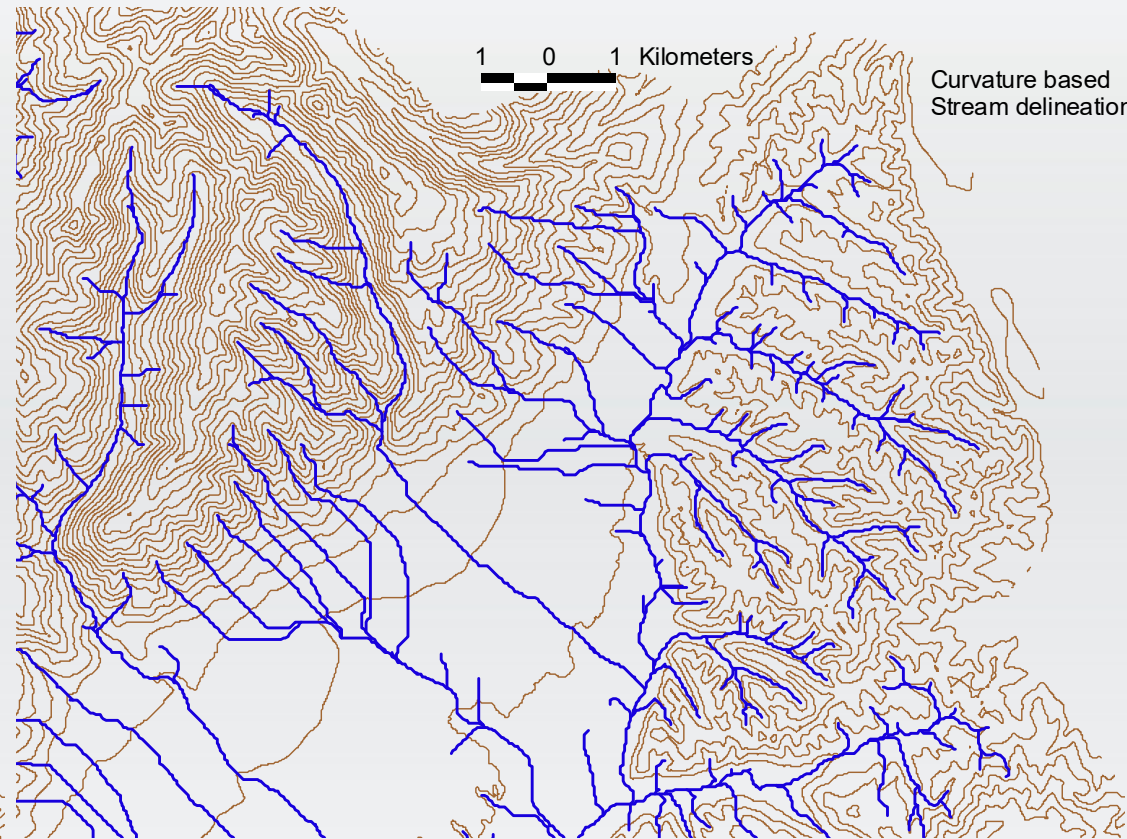
Weighted valley grid cell threshold	10	15	20	30
Drainage Density (km^{-1})	2.2	1.8	1.6	1.4
T for diff of first and higher order drop means	-4.1	-2.2	-1.3	-1.2



Channels mapped using contributing area and optimal threshold



Channels mapped using valley grid cells as weight and optimal threshold



Summary so far

- A simple grid DEM is the starting point for watershed delineation that produces structured information on watersheds and connected stream networks useful for hydrologic modeling
 - Knowledge from geomorphology offers an approach towards objective selection of channel network delineation threshold
 - Curvature and valley cell identification adapt for spatially variable drainage density and topographic texture
 - Watershed delineation based on D8
-
- Results are limited by being scale dependent and objective approach to finding a channelization threshold does not always work well
 - Opportunities to extend using other valley classification approaches (some seen at this conference)

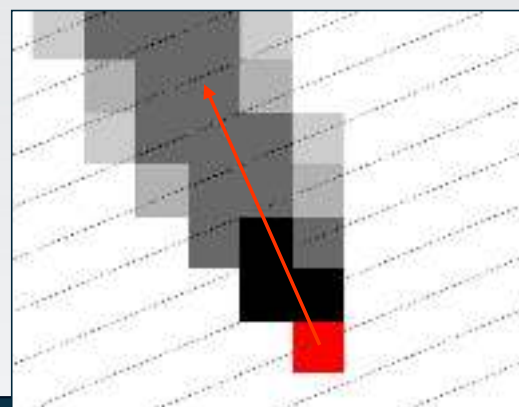
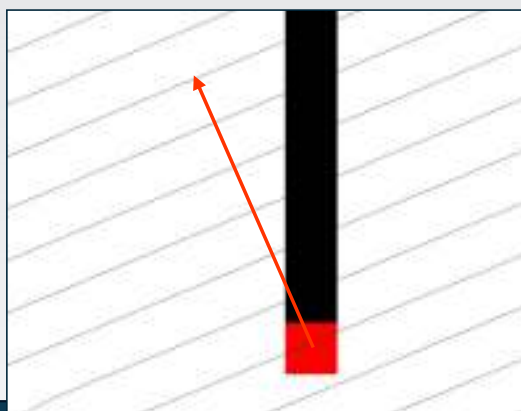
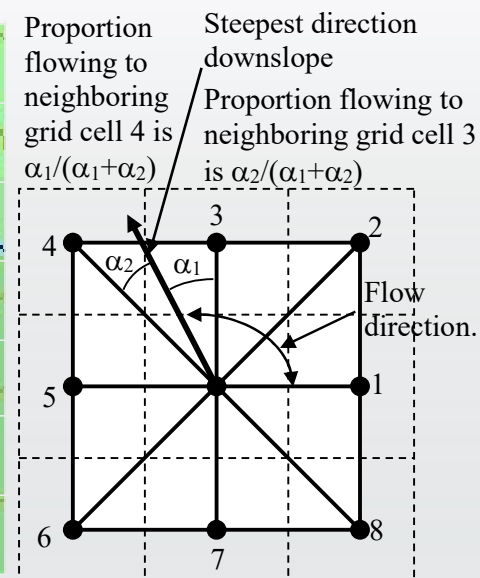
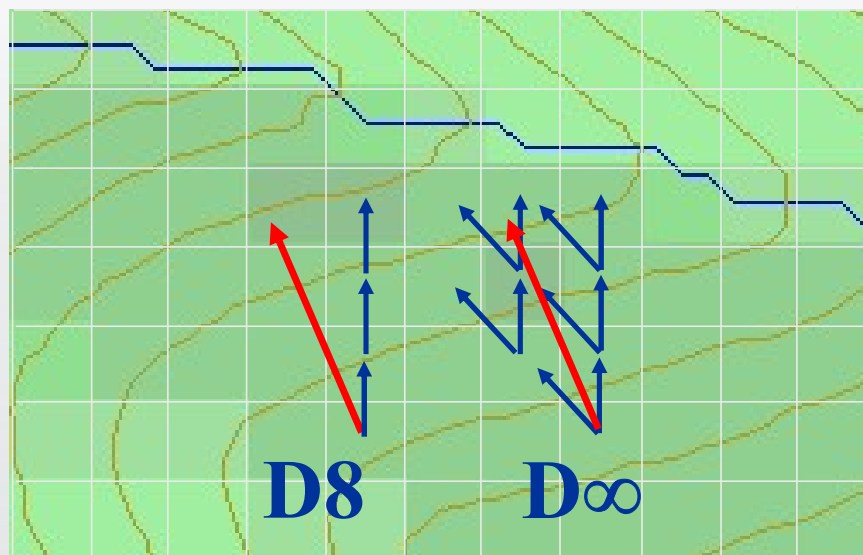


Representation of Flow Field

Steepest
single
direction

48	52
56	67

$$\frac{67 - 52}{30} = 0.50$$



Tarboton, D. G., (1997), "A New Method for the Determination of Flow Directions and Contributing Areas in Grid Digital Elevation Models," Water Resources Research, 33(2): 309-319, <http://doi.org/10.1029/96WR03137>.



Contributing Area

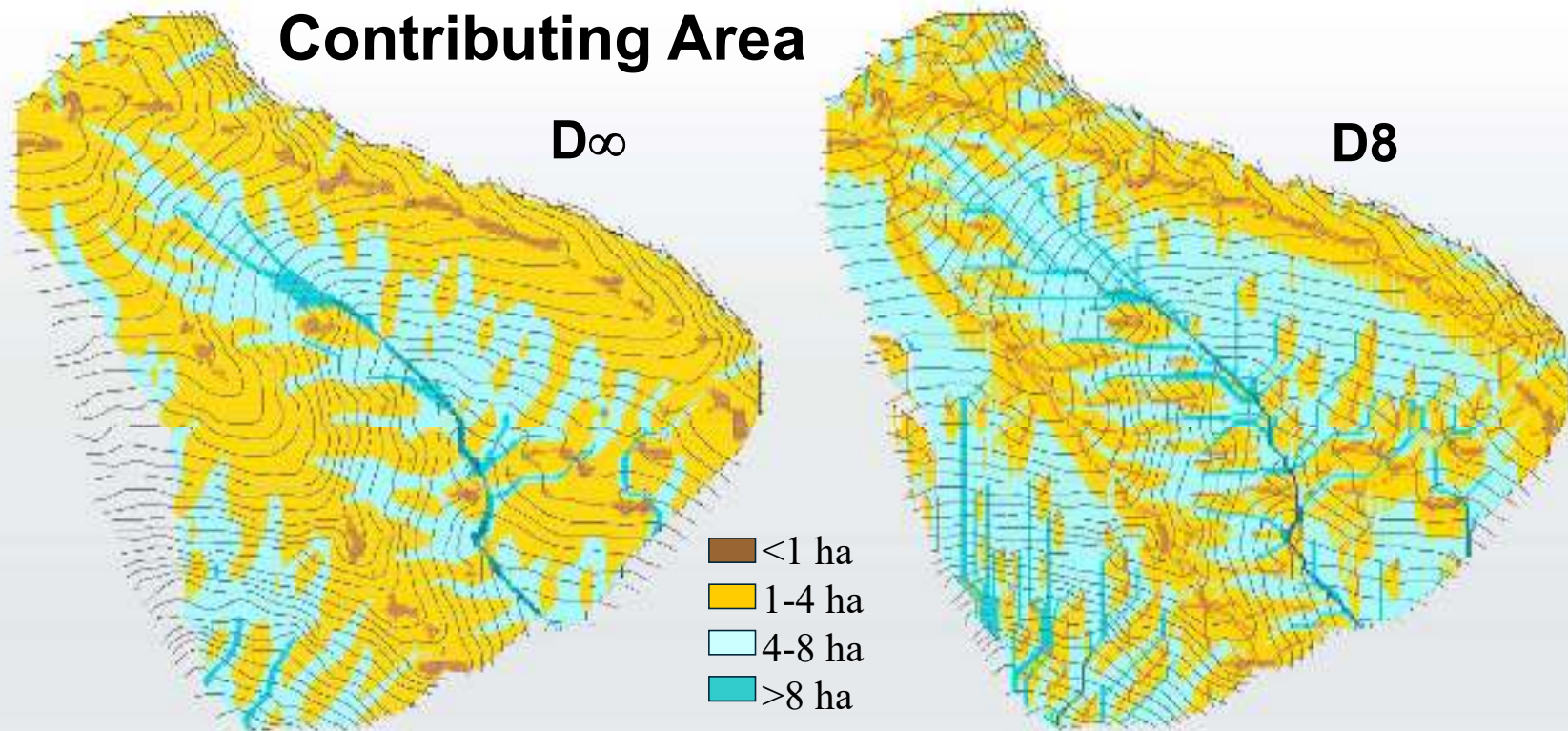
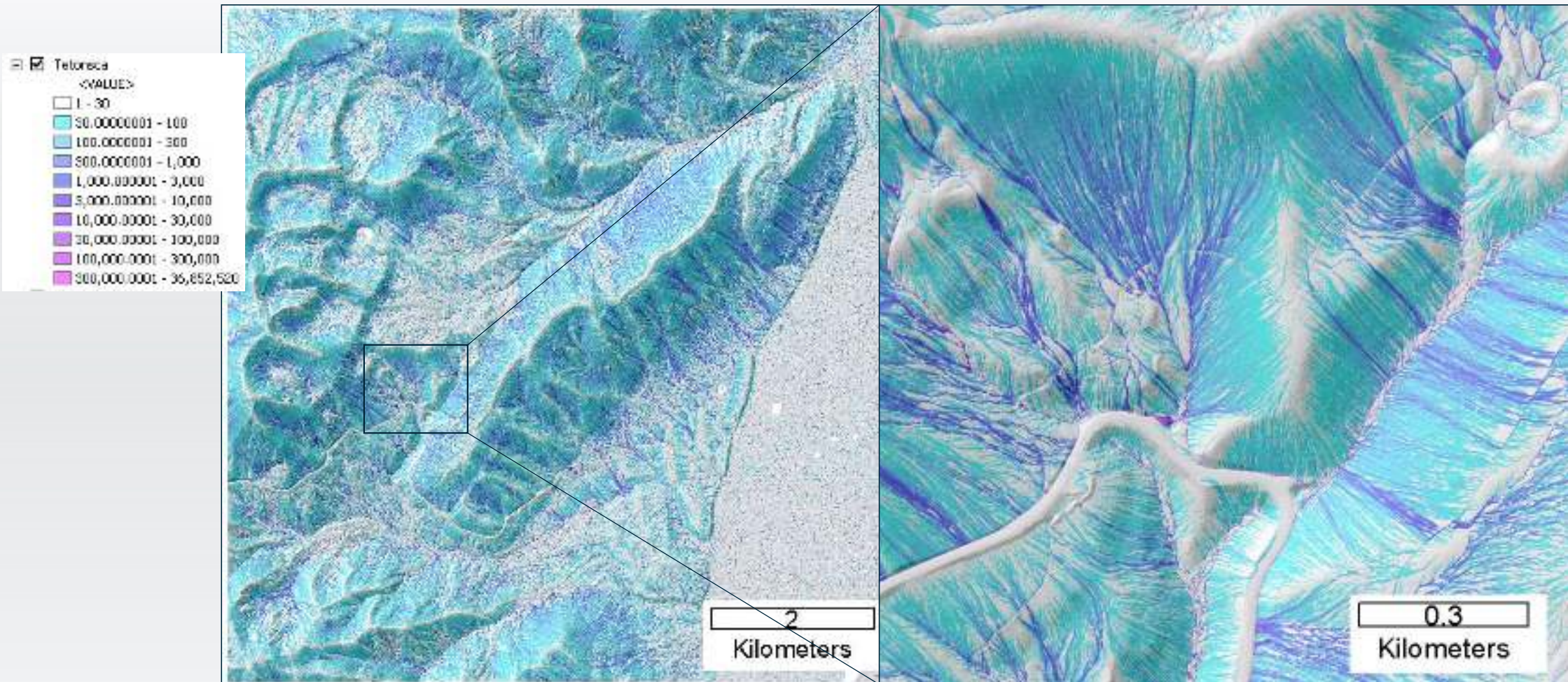


Table 2. Differences Between Theoretical and DEM-Computed Upslope Area for Test Examples Expressed in Terms of the Mean Error and Mean Square Error

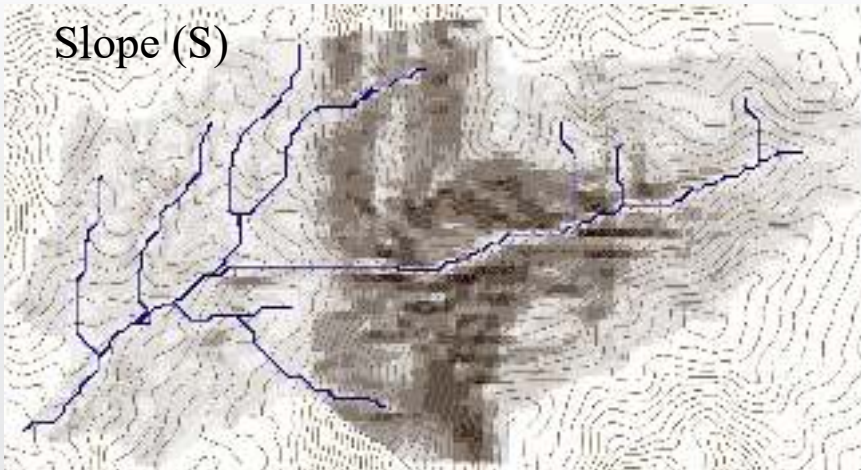
	Outward Cone		Inward Cone		Plane	
	Bias Mean $(A - \hat{A})$	MSE Mean $((A - \hat{A})^2)$	Bias Mean $(A - \hat{A})$	MSE Mean $((A - \hat{A})^2)$	Bias Mean $(A - \hat{A})$	MSE Mean $((A - \hat{A})^2)$
D8	-0.13	2.13	1.76	118.88	-0.17	0.065
MS	-0.81	0.69	-1.07	5.70	-1.37	2.065
Lee's [1992] method	-1.29	2.41	-4.05	44.00	-2.57	7.912
DEMON	-0.37	0.17	-0.37	19.23	-0.40	0.161
D ∞	-0.13	0.20	1.87	30.58	-0.17	0.065



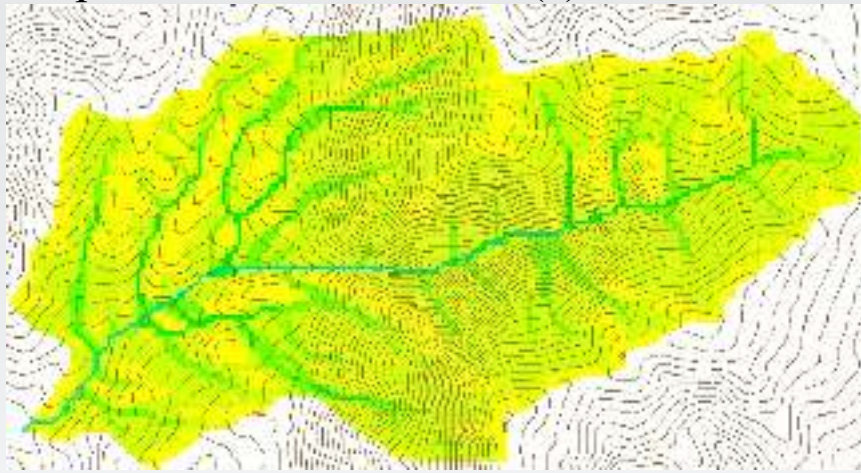
Contributing area from D-Infinity



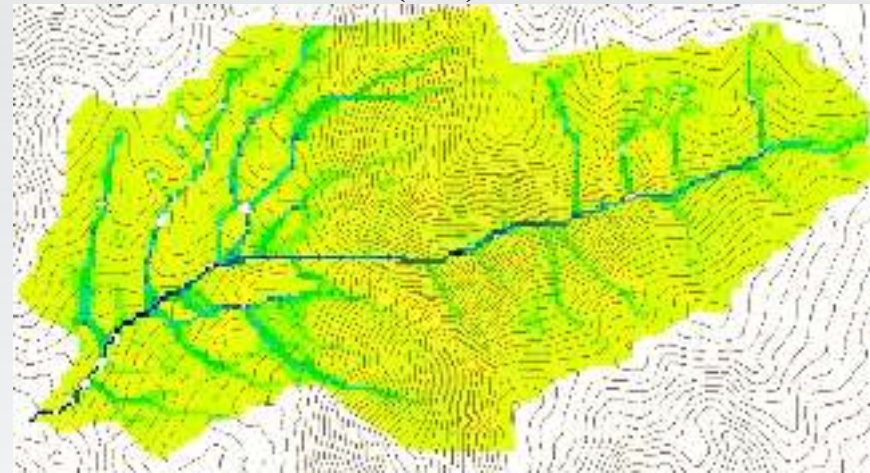
Wetness Index



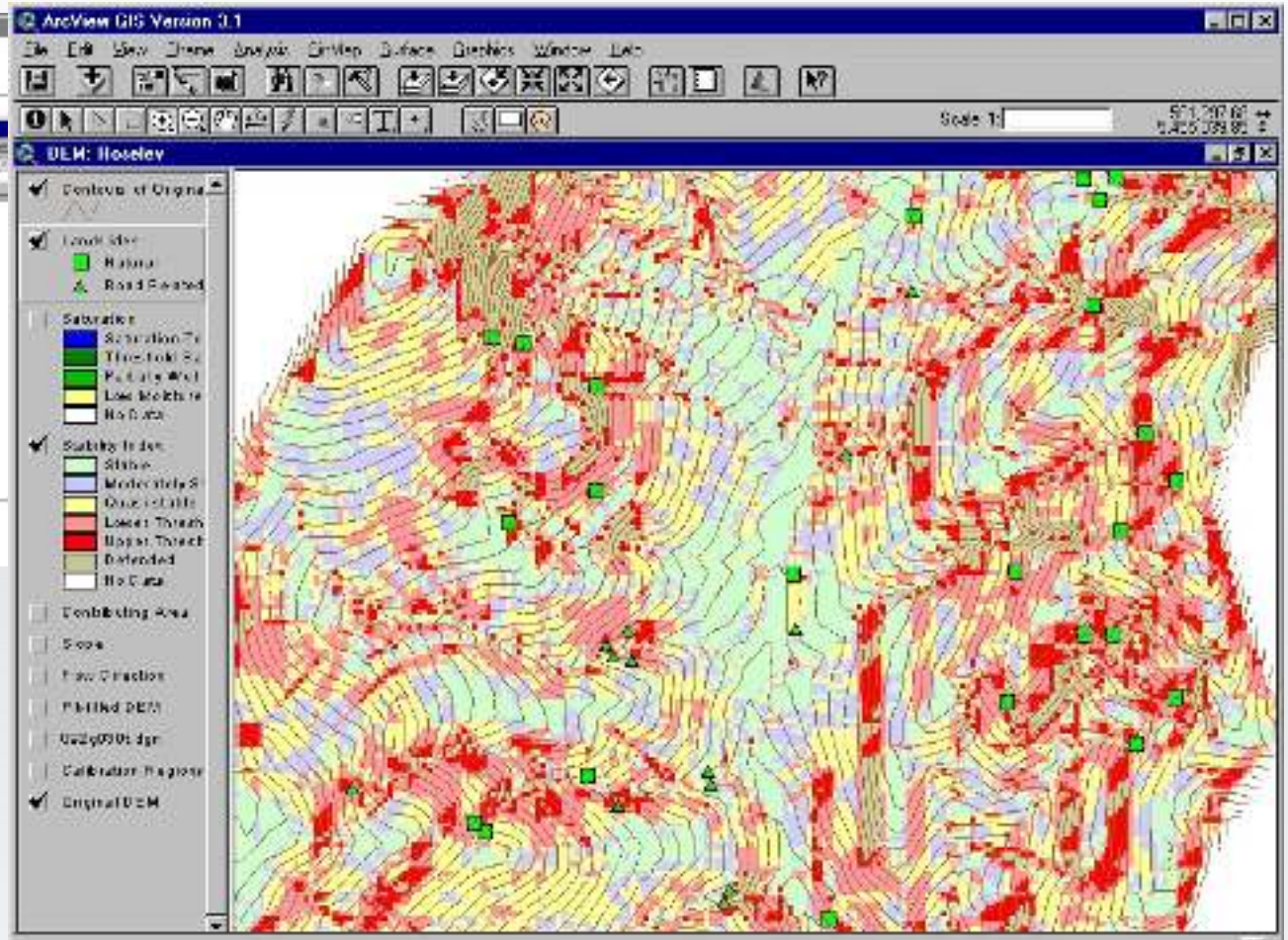
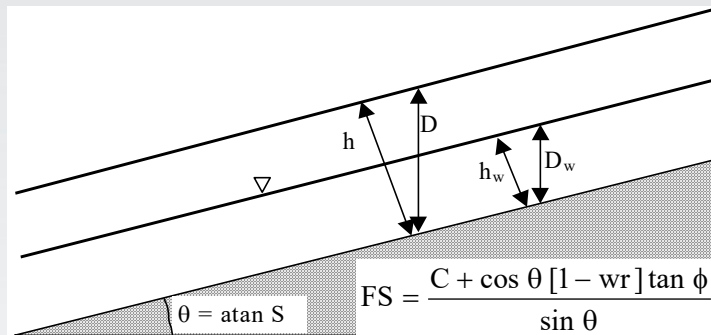
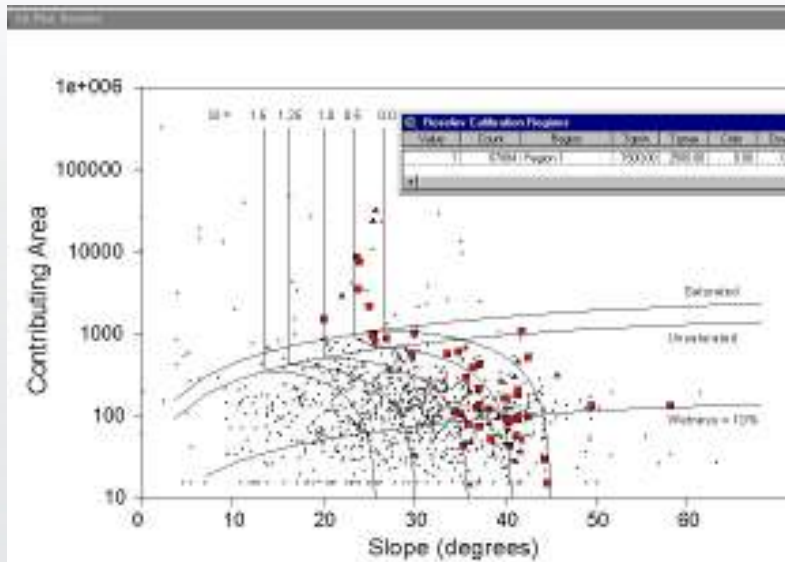
Specific Catchment Area (a)



Wetness Index $\ln(a/S)$



Terrain Stability Mapping



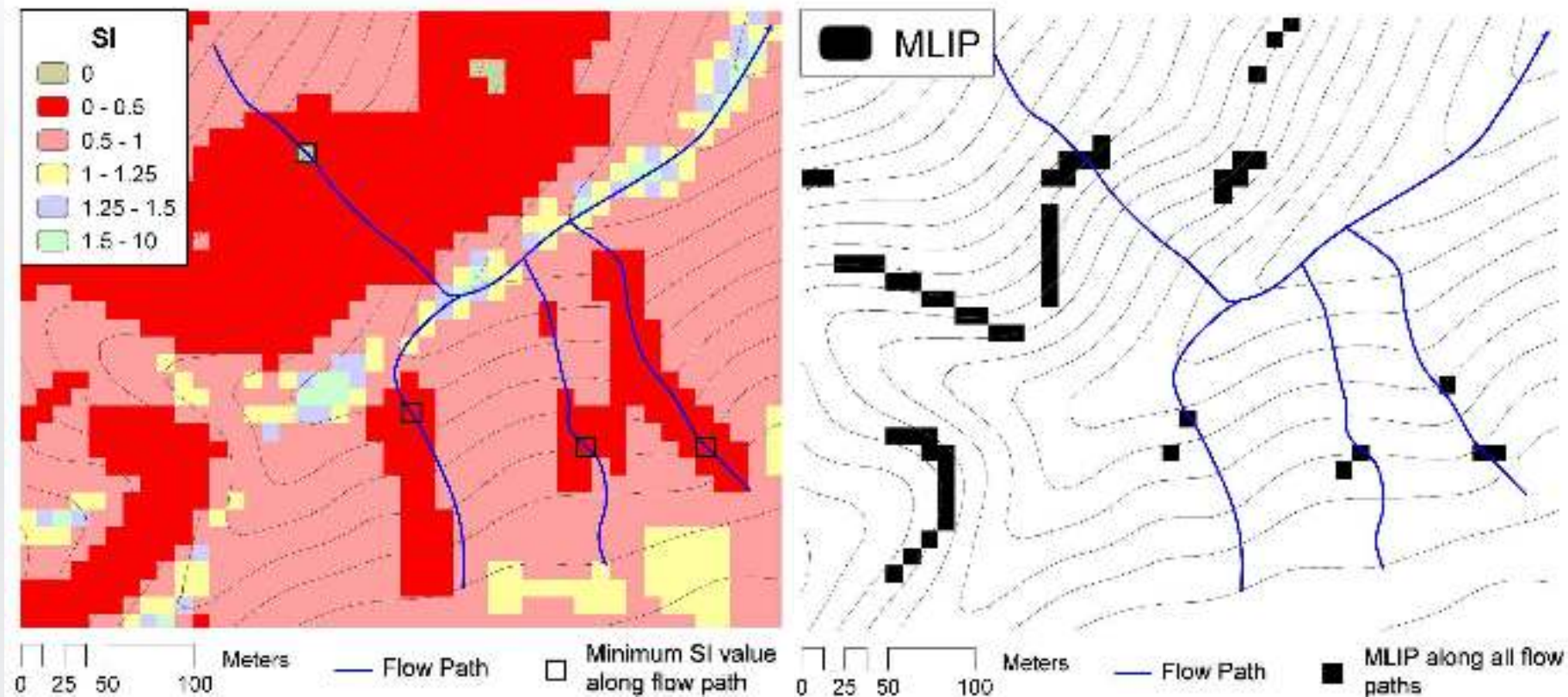
SINMAP: <https://hydrology.usu.edu/sinmap2/>



Utah Water Research Laboratory
Utah State University

<http://hydrology.usu.edu>

Most Likely Landslide Initiation Points



The location of the lowest SI value along a flow path

Tarolli, P. and D. G. Tarboton, (2006), HESS <https://doi.org/10.5194/hess-10-663-2006>



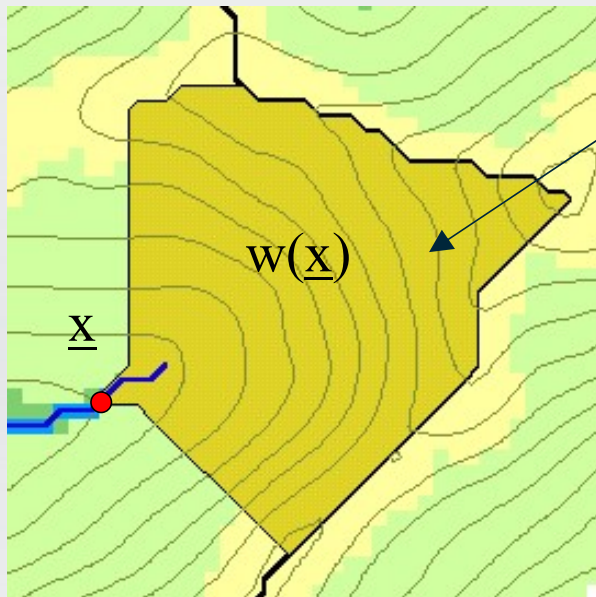
Other contributions, considerations and outstanding questions

- Seibert and McGlynn, 2007, MD ∞ , <http://dx.doi.org/10.1029/2006WR005128>
- Gallant and Hutchinson, 2011, differential equation approach, <http://doi.org/10.1029/2009wr008540>
- Gruber, S. and S. Peckham, 2009, [https://doi.org/10.1016/S0166-2481\(08\)00007-X](https://doi.org/10.1016/S0166-2481(08)00007-X) (Mass flux method)
- Lindsay (this conference) Minimal Dispersion Flow Algorithm
- Quinn et al., 1991; Fairfield and Leymarie, 1991

- Reduce angle/direction dispersion dependency
- Flow width. What cell dimension to use for specific catchment area calculation?
- Test cases beyond specific catchment area are needed



Flow accumulation calculation can be generalized to compute a broad class of terrain flow related quantities



$$A(\underline{x}) = \int_{CA} w(\underline{x}) d\underline{x}$$

The flow accumulation function takes as input a spatial field $w(\underline{x})$, and the topographic flow direction field and produces a field $A(\underline{x})$ representing the accumulation of $w(\underline{x})$ up to each point $\underline{x}$.

Numerical evaluation

$$A(i,j) = w(i,j)\Delta + \sum_{k \text{ contributing neighbors}} p_k A(i_k, j_k)$$

p_k is the proportion of flow from neighbor k contributing to the grid cell (i,j) .

$\sum p_k = 1$ is required to ensure 'conservation'.

Flow directions must not have loops.

Generalization to Flow Algebra

Replace

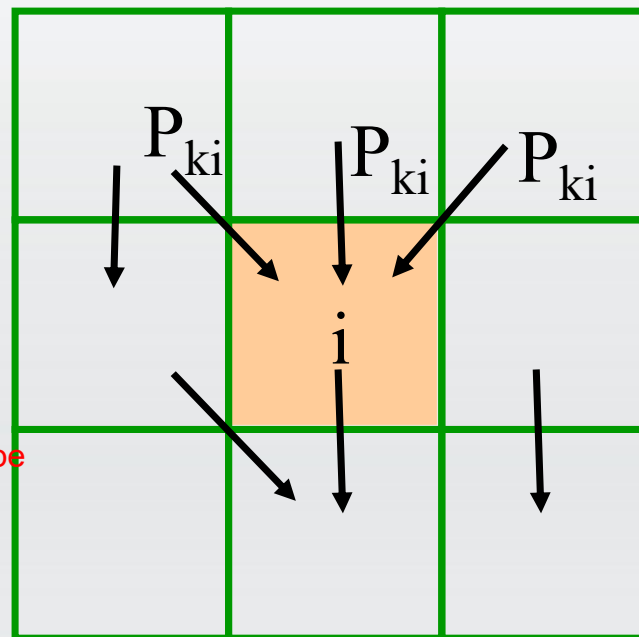
$$A_i = w_i \Delta + \sum_{\text{contributing } k} P_{ki} A_k$$

by general function

$$\theta_i = F(\gamma_i, P_{ki}, \theta_k, \gamma_k)$$

Flow direction proportions Input Results from upslope

Output Result, Quantity being calculated Other inputs at cell and upslope

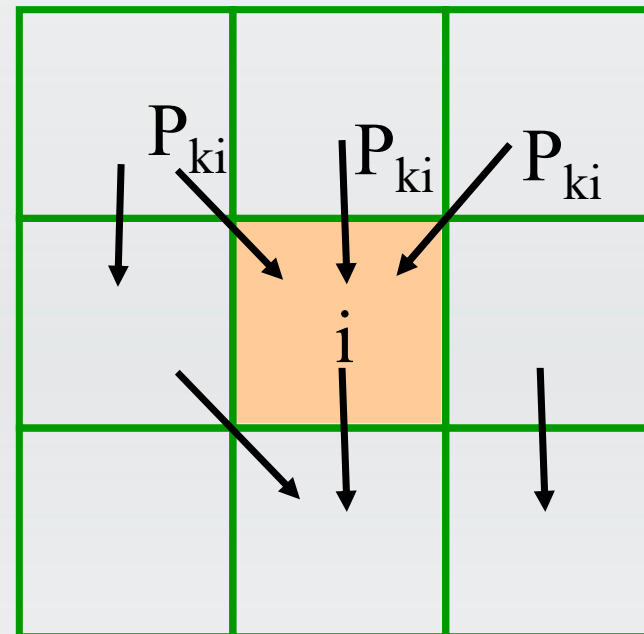


Tarboton, D. G. and M. E. Baker, (2008), "Towards an Algebra for Terrain-Based Flow Analysis," in Representing, Modeling and Visualizing the Natural Environment: Innovations in GIS 12, Edited by N. J. Mount, G. L. Harvey, P. Aplin and G. Priestnall, CRC Press, Florida, p.496, <http://dx.doi.org/10.1201/9781420055504.ch12>.



General Pseudocode **Upstream** Flow Algebra Evaluation

Global $\underline{P}$, $\underline{\gamma}$, $\underline{\theta}$
FlowAlgebra(i)
for all k neighbors of i
 if $P_{ki} > 0$
 FlowAlgebra(k)
next k
 $\underline{\theta}_i = \text{FA}(\underline{\gamma}_i, \underline{P}_{ki}, \underline{\theta}_k, \underline{\gamma}_k)$
return

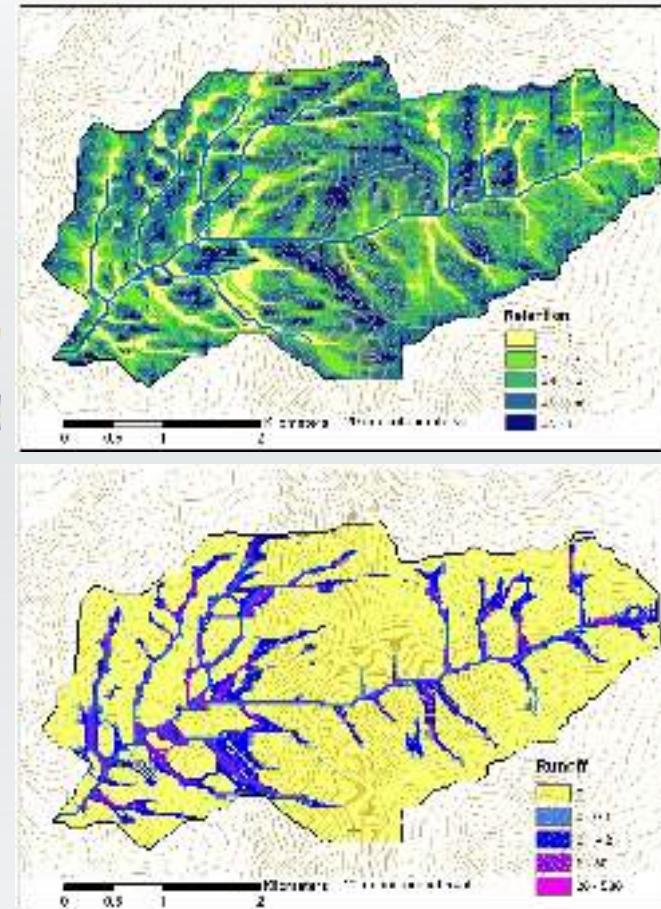
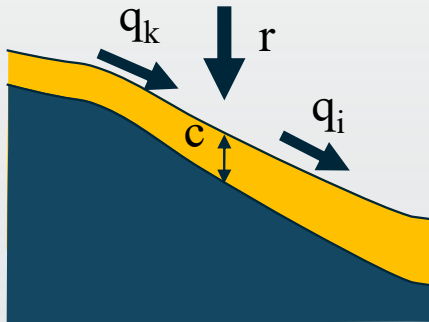


Example: Retention limited runoff generation with run-on

Global $\underline{P}$, $(\underline{r}, \underline{c})$, $\underline{q}$
FlowAlgebra(i)
for all k neighbors of i
if $P_{ki} > 0$
 FlowAlgebra(k)
next k

$$q_i = \max\left(\sum_{\{k:P_{ki}>0\}} P_{ki} q_k + r_i - c_i, 0\right)$$

return



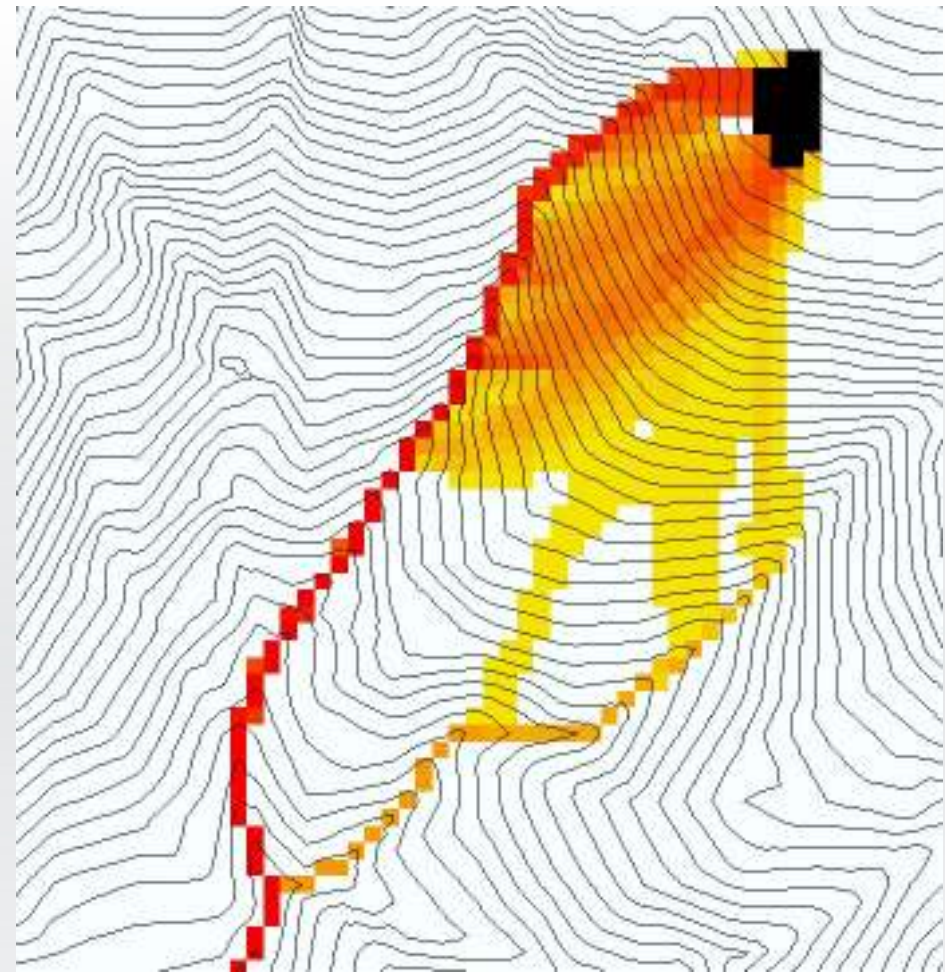
Decaying Accumulation

Mass loading field $m(x)$ assumed to move with the flow field but is subject to first order decay

$$DA(i,j) = m(i,j)\Delta^2 + \sum_{k \text{ contributing neighbors}} p_k d(i_k, j_k) DA(i_k, j_k)$$

$d(i,j)$ is a decay multiplier giving the reduction in mass in moving from one grid cell to the next. $d(i,j)$ may be related to travel time, e.g. as

$d(i,j) = \exp(-\lambda t(i,j))$ where λ is a first order decay parameter.



Useful for a tracking contaminant or compound subject to decay or attenuation



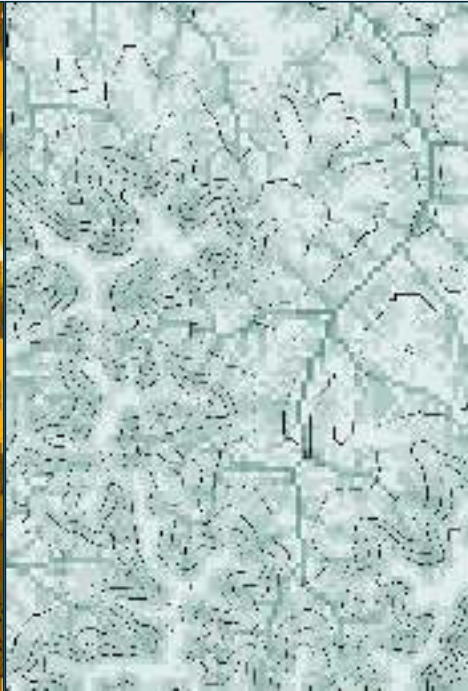
Transport limited accumulation

Supply



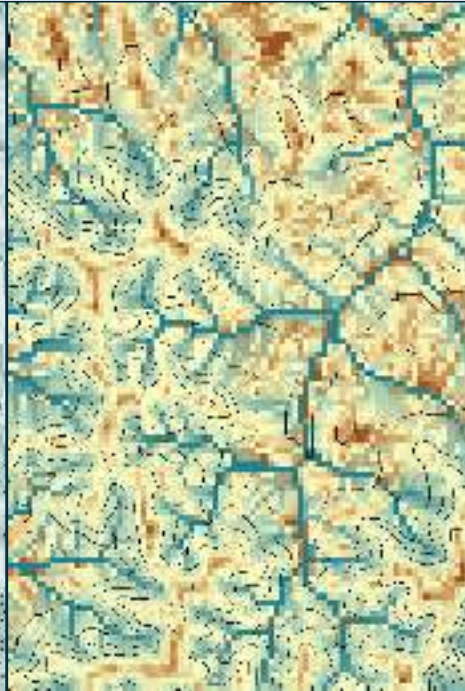
S

Capacity



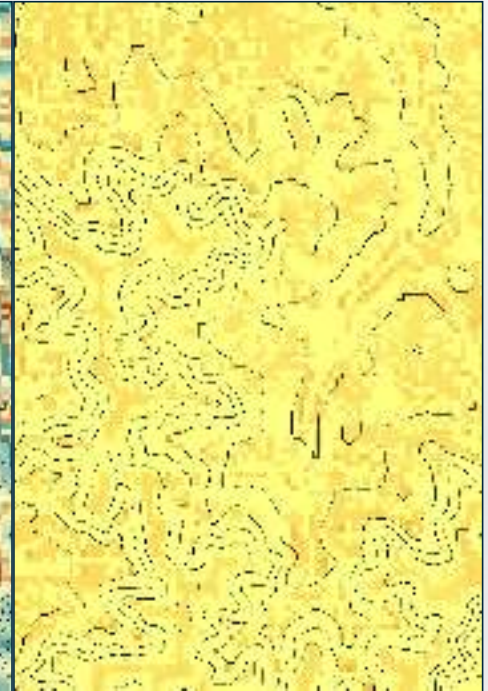
$$T_{cap} = \chi a^2 \tan(b)^2$$

Transport



$$T_{out} = \min\{S + \sum T_n, T_{cap}\}$$

Deposition



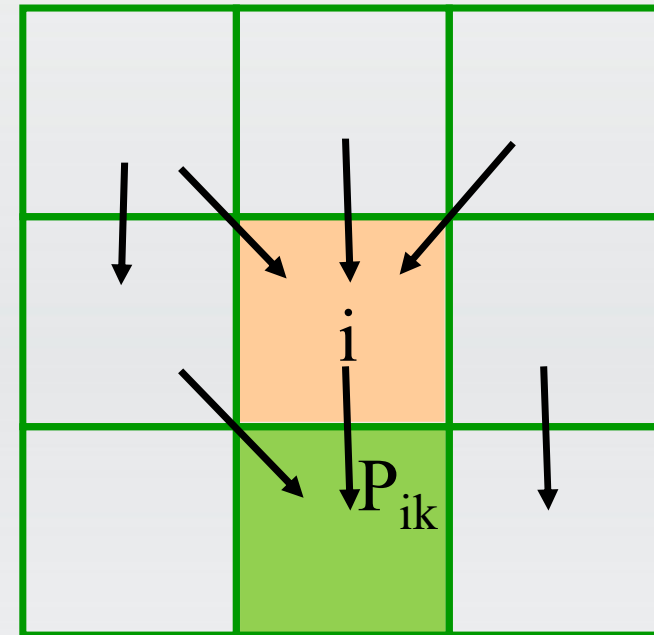
$$D = S + \sum T_n - T_{out}$$

Useful for modeling erosion and sediment delivery, the spatial dependence of sediment delivery ratio and contaminant that adheres to sediment

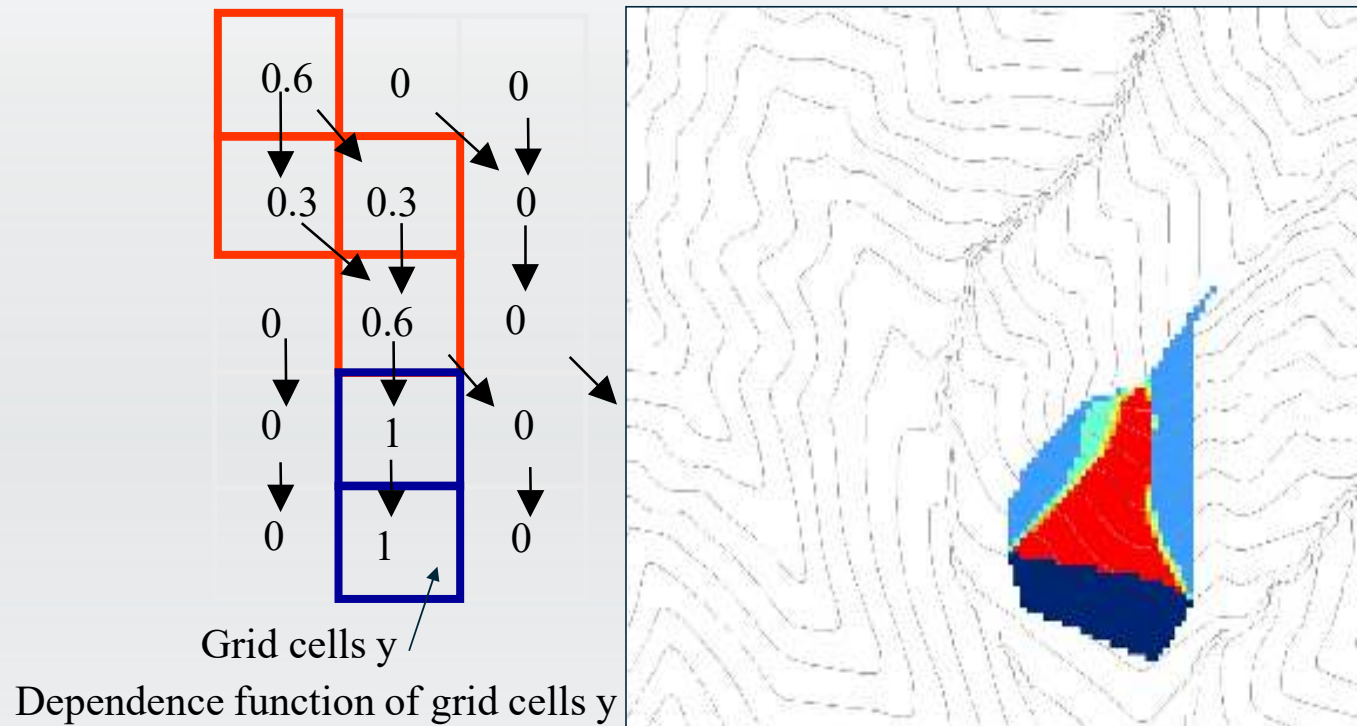


General Pseudocode **Downstream** Flow Algebra Evaluation

Global $\underline{P}$, γ , $\underline{\theta}$
FlowAlgebra(i)
for all k neighbors of i
 if $P_{ik} > 0$
 FlowAlgebra(k)
next k
 $\underline{\theta}_i = \text{FA}(\gamma_i, \underline{P}_{ik}, \underline{\theta}_k, \gamma_k)$
return



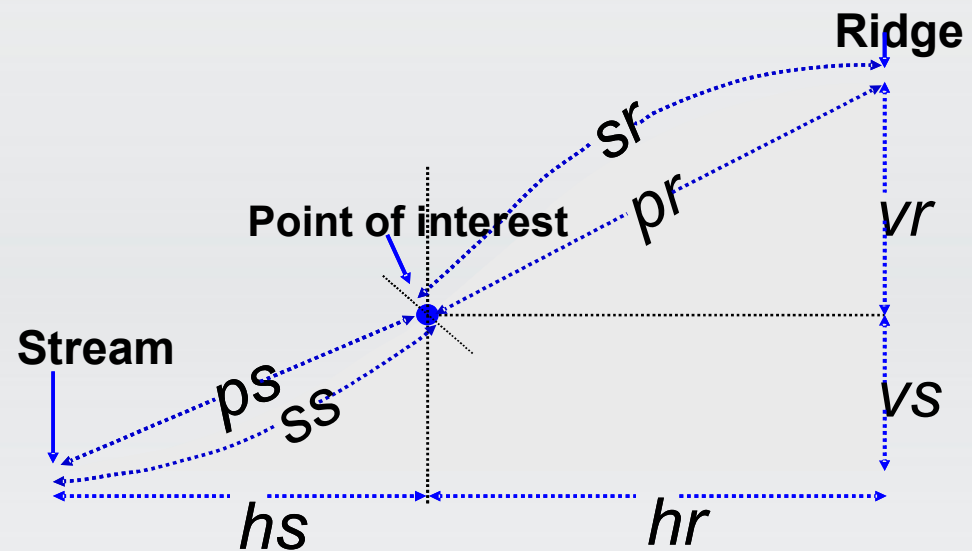
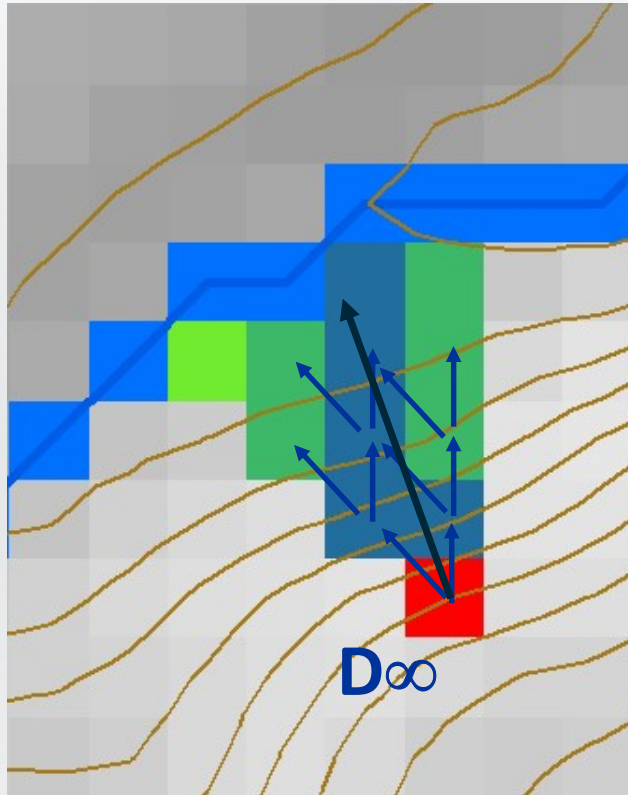
Dependence function. Quantifies the amount a point x contributes to the point or zone y.



Useful for example to track where a contaminant may come from



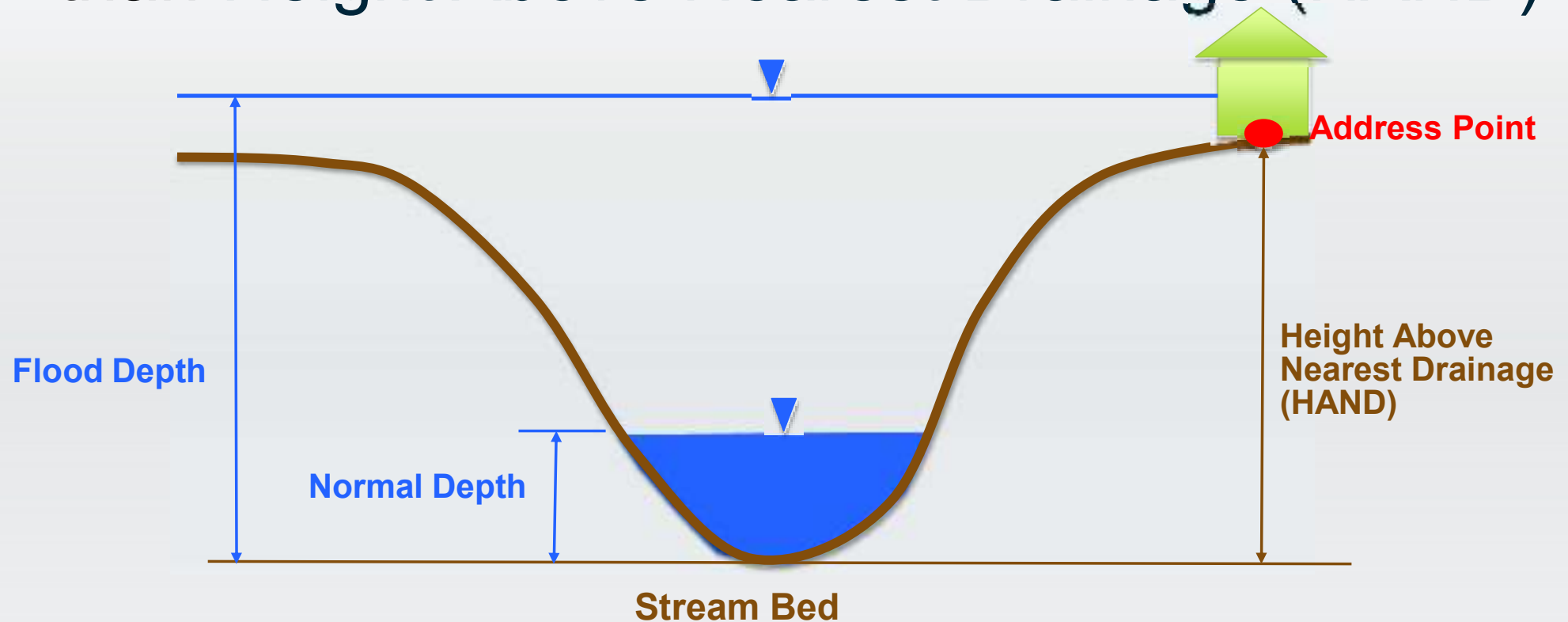
General Proximity Functions as weighted distance to a target set along Dinfinity Flow Directions



Tesfa, T. K., D. G. Tarboton, D. W. Watson, K. A. T. Schreuders, M. E. Baker and R. M. Wallace, (2011), "Extraction of hydrological proximity measures from DEMs using parallel processing," Environmental Modelling & Software, 26(12): 1696-1709, <http://dx.doi.org/10.1016/j.envsoft.2011.07.018>.



Flooding occurs when **Water Depth** is greater than **Height Above Nearest Drainage (HAND)**



May be used to determine flood risk

Slide from David Maidment



Utah Water Research Laboratory
Utah State University

<http://hydrology.usu.edu>

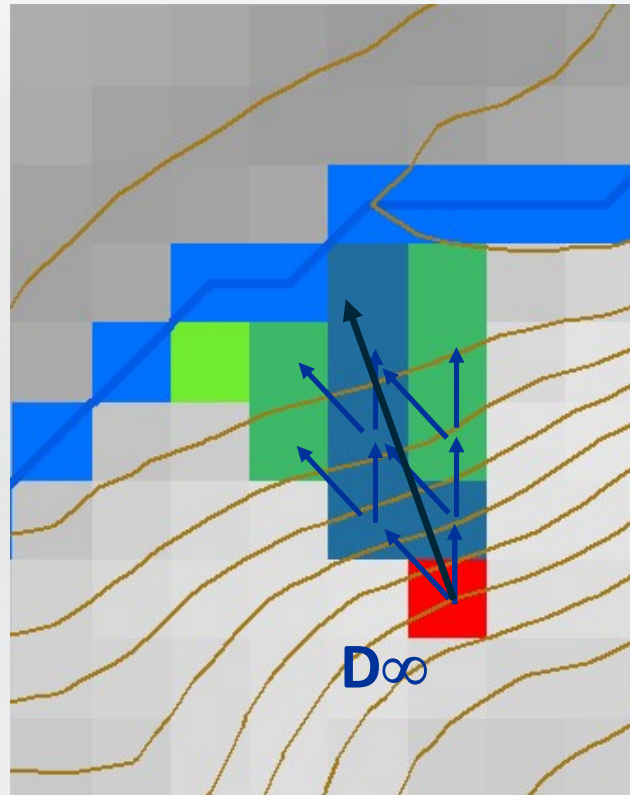
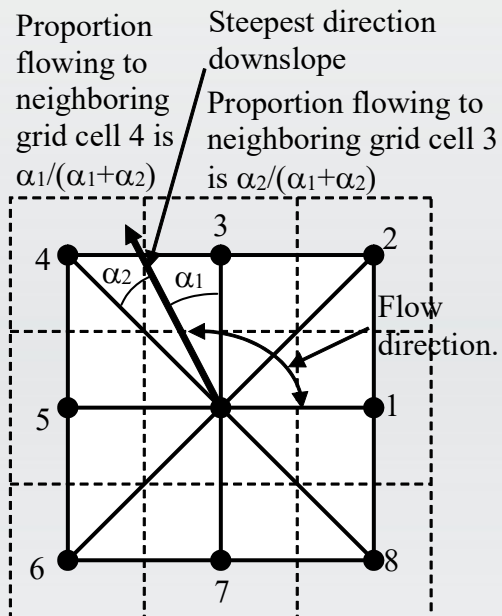
Literature on HAND

- Rodda, H. J. E., (2005), "The Development and Application of a Flood Risk Model for the Czech Republic," Natural Hazards, 36(1): 207-220, <http://dx.doi.org/10.1007/s11069-004-4549-4>.
- Rennó, C. D., A. D. Nobre, L. A. Cuartas, J. V. Soares, M. G. Hodnett, J. Tomasella and M. J. Waterloo, (2008), "HAND, a new terrain descriptor using SRTM-DEM: Mapping terra-firme rainforest environments in Amazonia," Remote Sensing of Environment, 112(9): 3469-3481, <http://doi.org/10.1016/j.rse.2008.03.018>.
- Nobre, A. D., L. A. Cuartas, M. Hodnett, C. D. Rennó, G. Rodrigues, A. Silveira, M. Waterloo and S. Saleska, (2011), "Height Above the Nearest Drainage – a hydrologically relevant new terrain model," Journal of Hydrology, 404(1–2): 13-29, <http://dx.doi.org/10.1016/j.jhydrol.2011.03.051>.
- Nobre, A. D., L. A. Cuartas, M. R. Momo, D. L. Severo, A. Pinheiro and C. A. Nobre, (2016), "HAND contour: a new proxy predictor of inundation extent," Hydrological Processes, 30(2): 320-333, <http://dx.doi.org/10.1002/hyp.10581>.

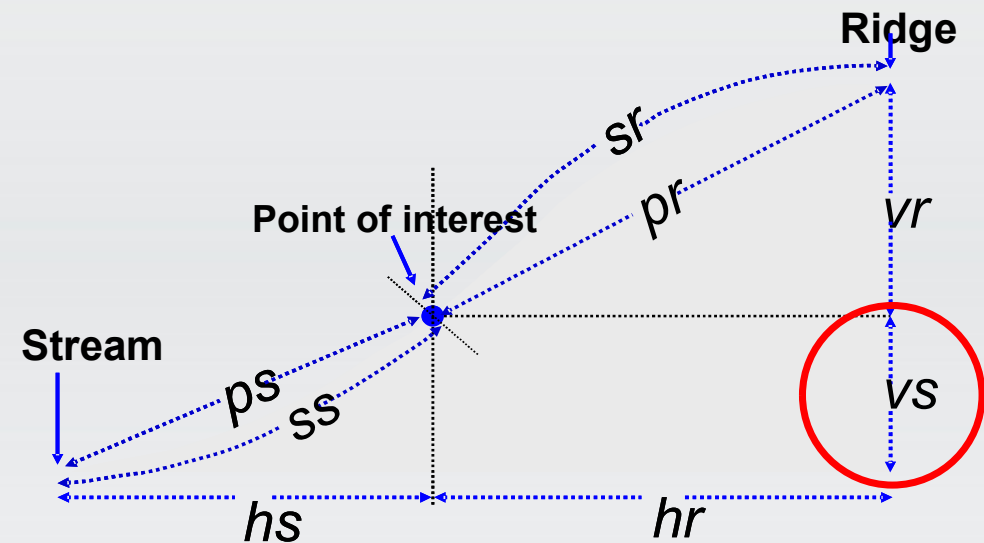


Height above Nearest Drainage (HAND) evaluated using TauDEM Dinfinity vertical distance down function

D_{∞}



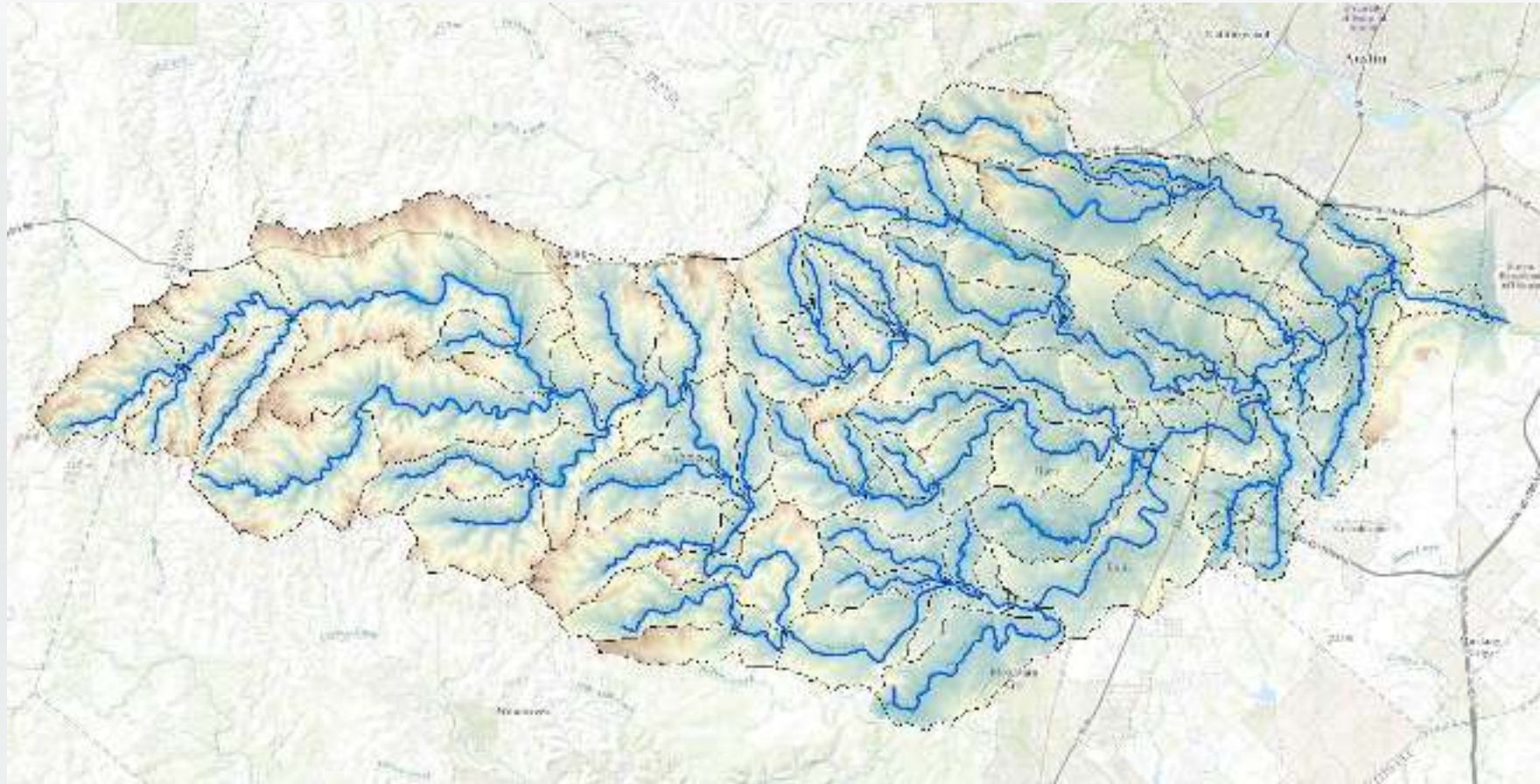
A special case of a general hydrological proximity measure



Vertical distance to stream evaluated as weighted average over multiple flow paths.
This results in a “smooth” height above nearest drainage layer



HAND from TauDEM for Onion Creek near Austin Texas



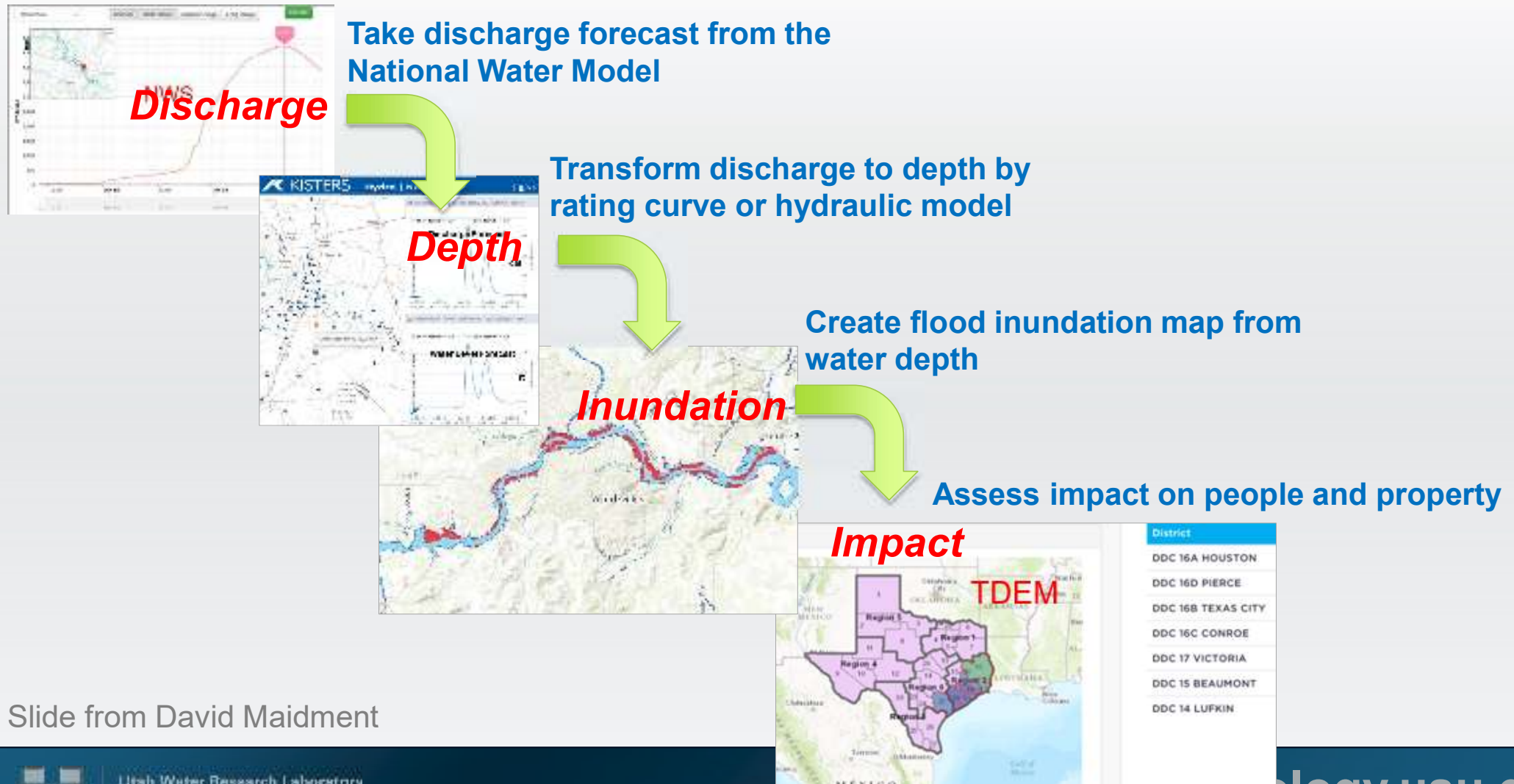
1/3 arc second NED DEM



Utah Water Research Laboratory
UtahStateUniversity

<http://hydrology.usu.edu>

Flood Inundation Mapping from National Water Model Discharges



Slide from David Maidment



Utah Water Research Laboratory
Utah State University

<http://hydrology.usu.edu>

Implementation

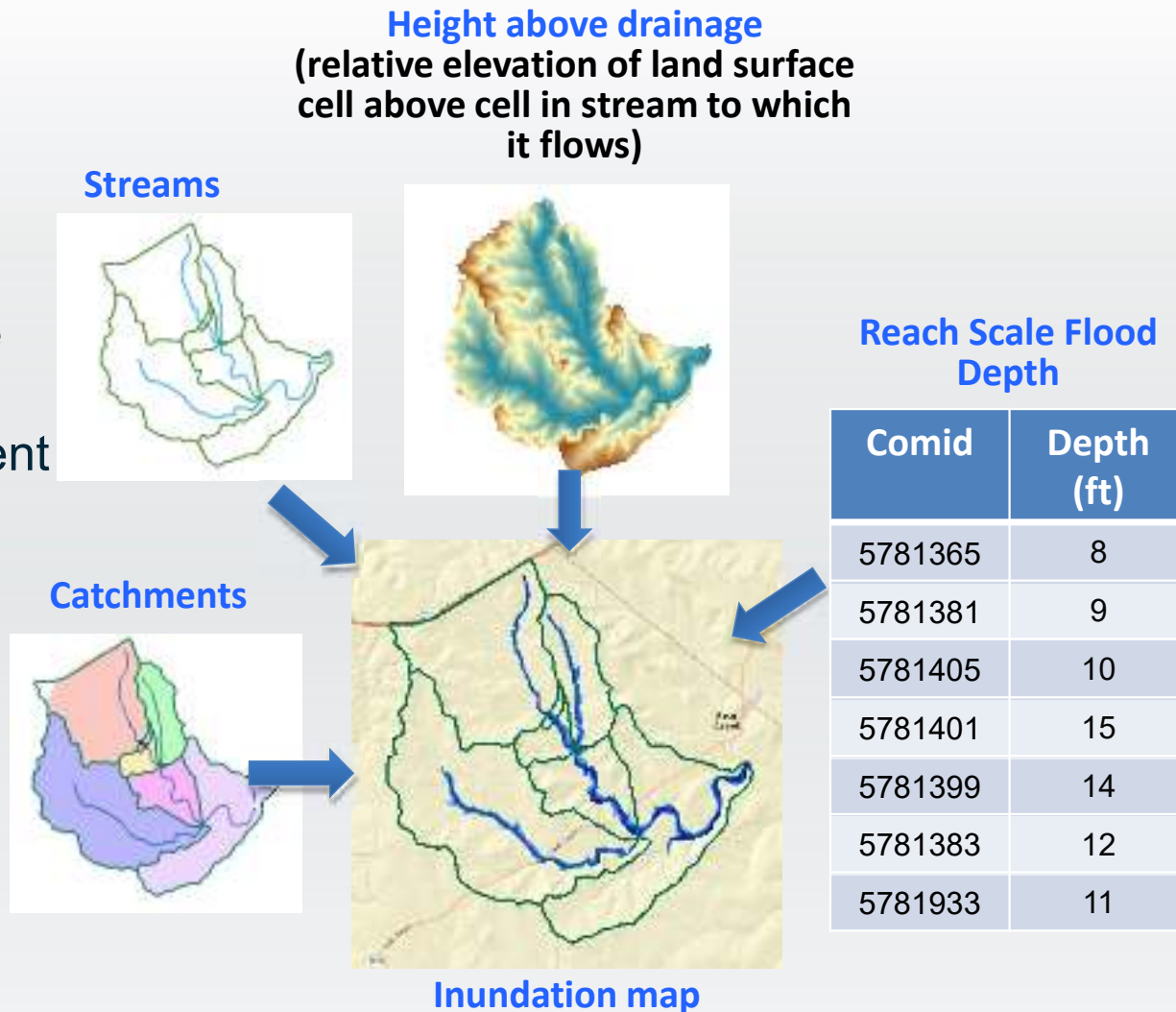
- Each stream reach has a discharge
- Each stream reach has a water depth calculated from discharge, h_w
- Each stream reach has a unique ID
- Each catchment has a unique ID for the reach to which it drains
- Each grid cell has the ID of the catchment it is in (reach it drains to) and height above the nearest drainage, h_d
- Evaluate inundation using

if($h_w(id) > h_d(id)$)

Inundation depth = $h_w(id) - h_d(id)$

else

Inundation depth = 0



Using HAND to determine reach average hydraulic properties

For each Catchment
For each height h
Identify cells where HAND $h_d < h$

Single Cell Plan Area $A_c = dx * dy$

Surface area $A_s = \sum A_c$

Bed Area $A_b = \sum A_c \sqrt{1 + slp^2}$

Approximates each cell as sloping plane

Volume $V = \sum A_c (h - h_d)$

$A = \frac{V}{L}$ Cross Section Area

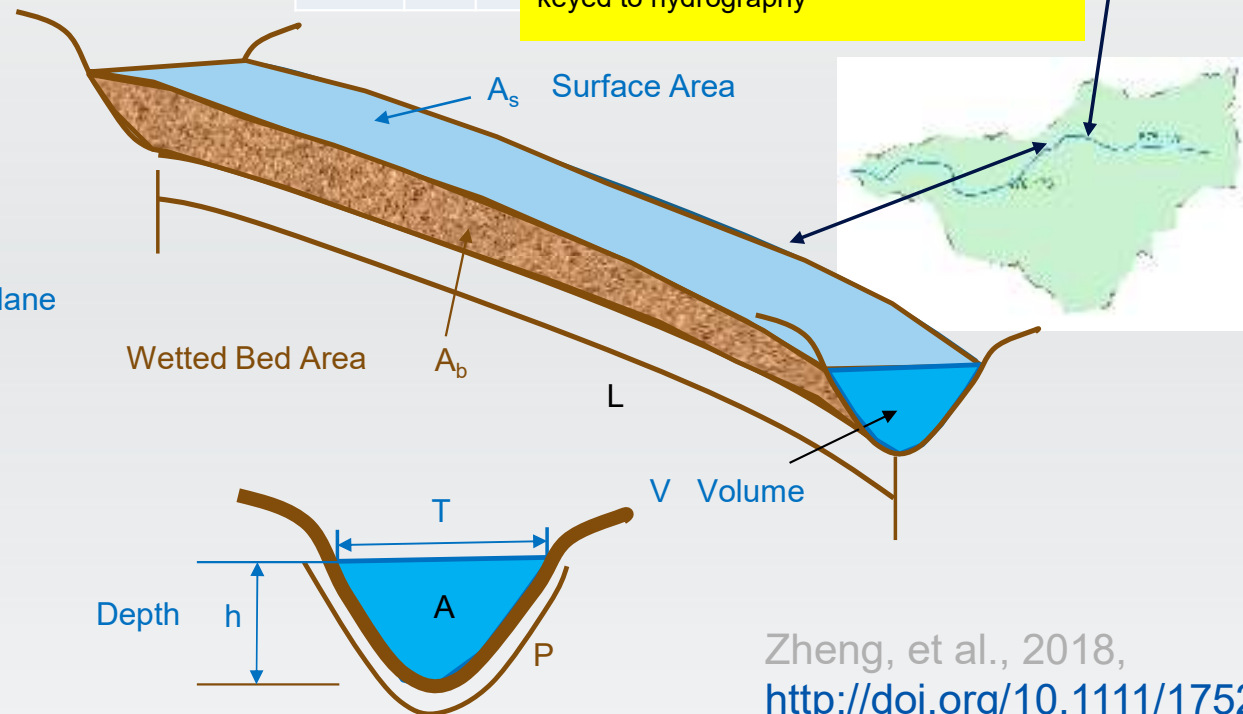
$P = \frac{A_b}{L}$ Wetted Perimeter

$T = \frac{A_s}{L}$ Top Width

$R = \frac{A}{P}$ Hydraulic Radius

Comid	H	A	R	P	T	V	Ab	As
5781175	3							
5781175	4							

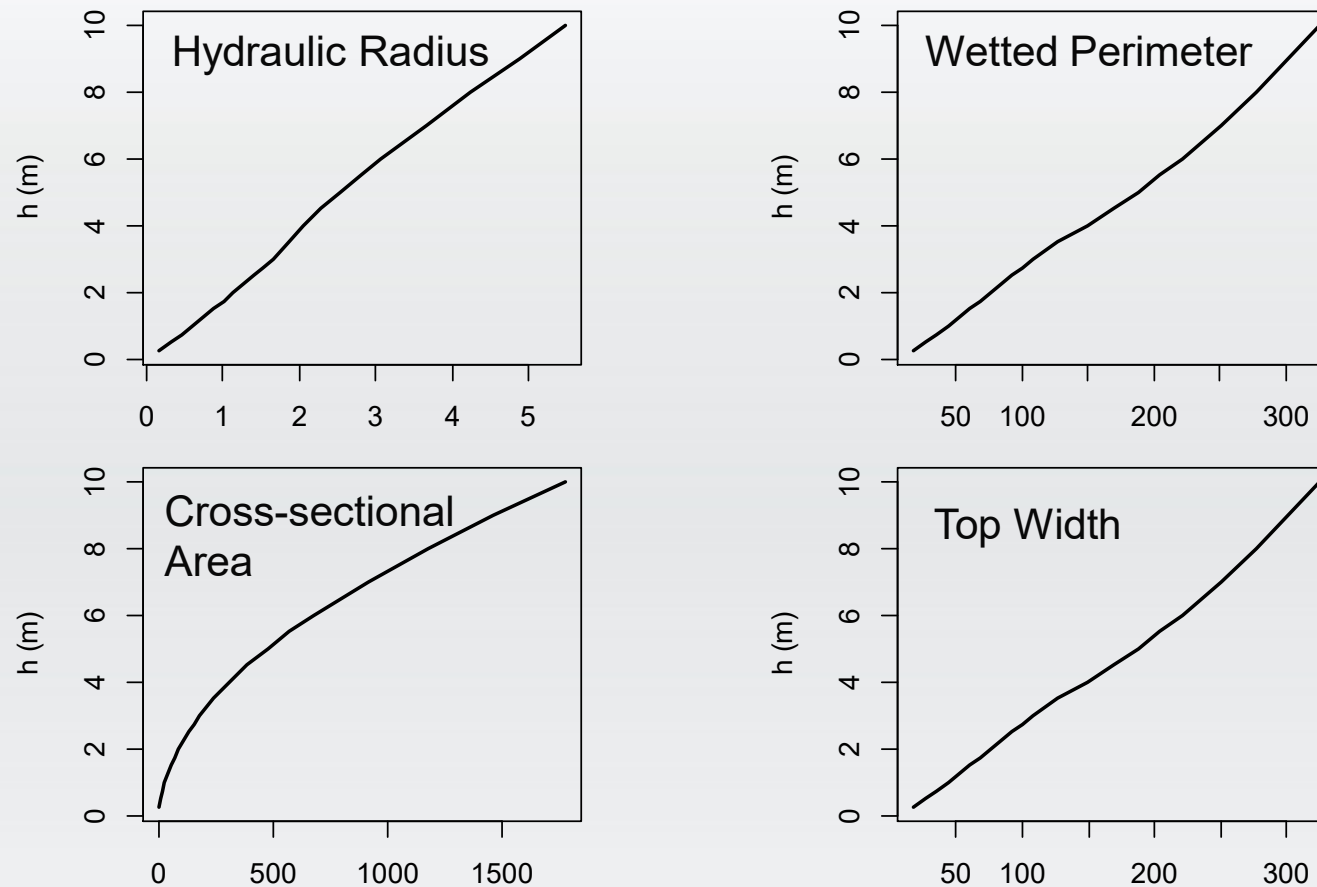
A table with reach hydraulic parameters keyed to hydrography



Zheng, et al., 2018,
<http://doi.org/10.1111/1752-1688.12661>.



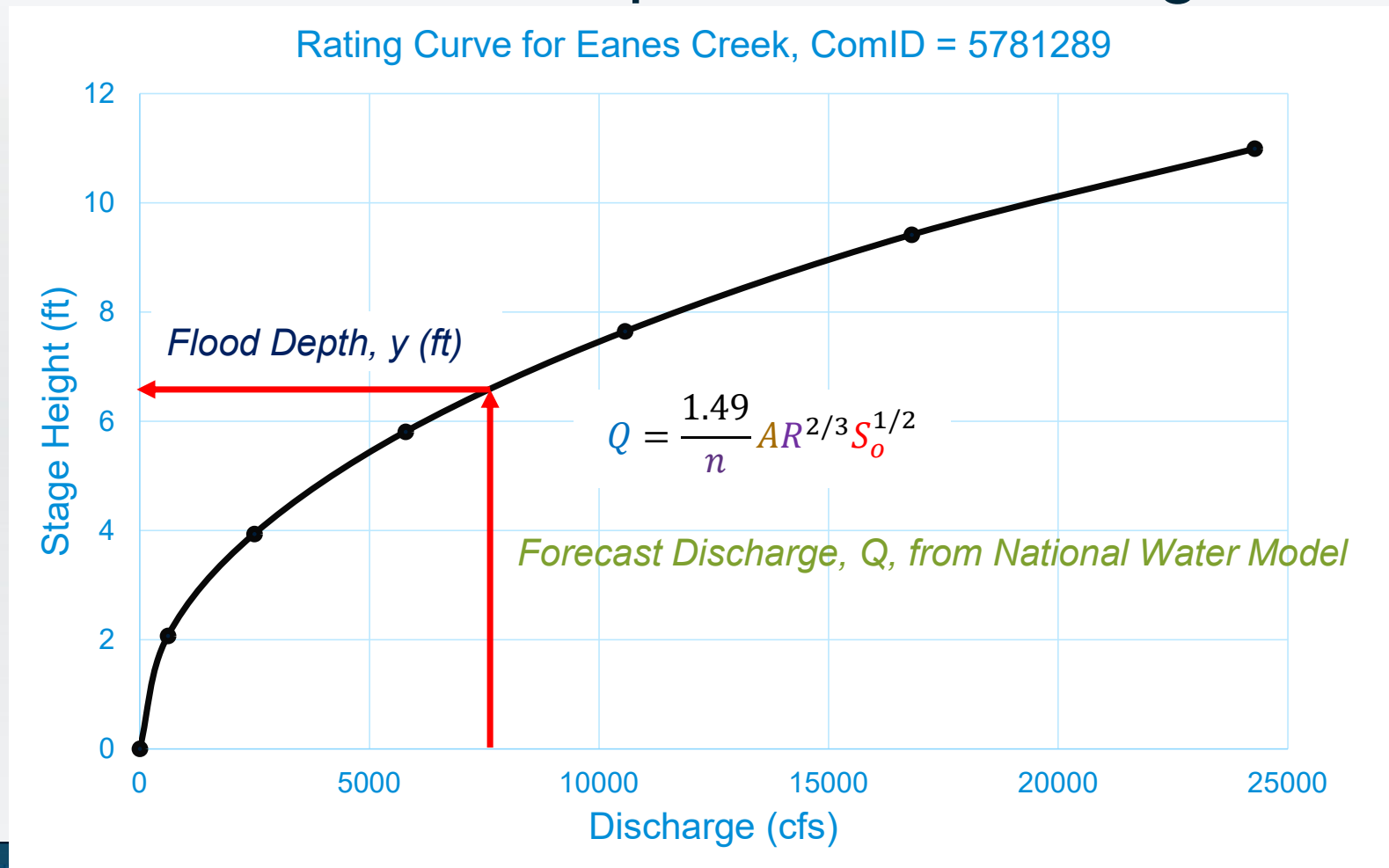
Terrain Approximated Reach Average Hydraulic Properties



Used for input to reach scale hydraulic model



Combined with roughness estimate to determine rating curve and depth from discharge



Slide from David
Maidment



The challenge of increasing Digital Elevation Model (DEM) resolution demands a parallelization approach

1980's DMA 90 m
 10^2 cells/km²

1990's USGS DEM 30
 10^3 cells/km²

2000's NED 10 m
 10^4 cells/km²

2010's LIDAR ~1 m
 10^6 cells/km²



e.g. 50,000 km²
Watershed
27 MB

240 MB

2 GB

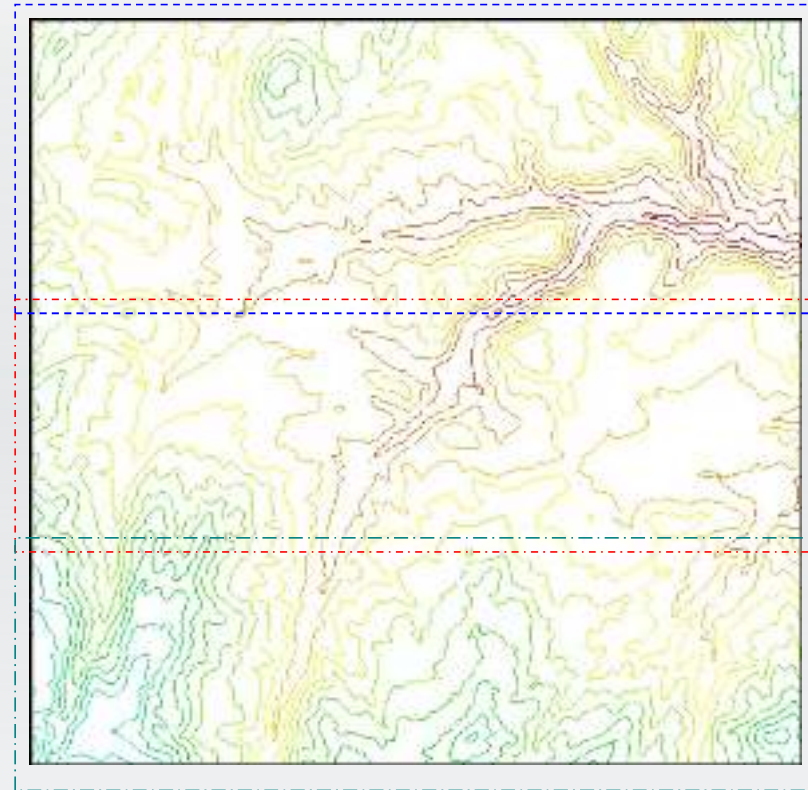
200 GB

A factor 10,000
increase in 30 years

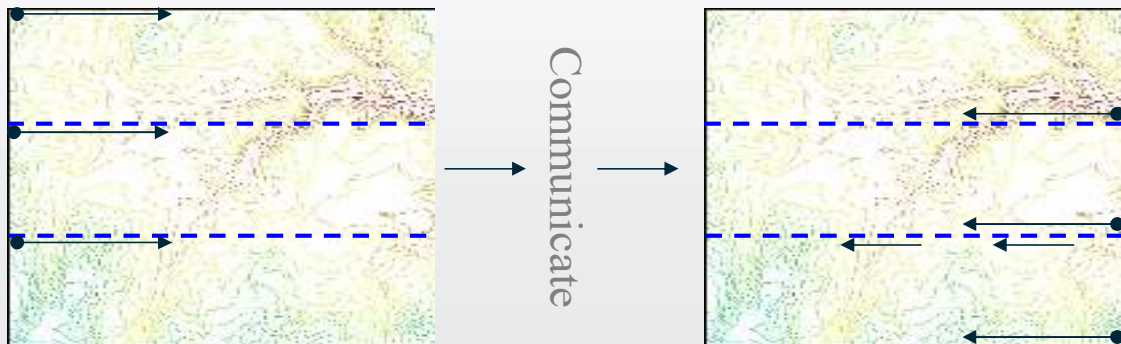


TauDEM Parallel Programming Approach

- MPI, distributed memory paradigm
- Row oriented slices
- Each process includes one buffer row on either side
- Each process does not change buffer row
- Improved runtime efficiency
- Capability to run larger problems



Parallel Pit Filling based on Planchon and Darboux approach



```

Initialize( Z,F)
Do
  for all grid cells i
    if  $Z(i) > n$ 
       $F(i) \leftarrow Z(i)$ 
    Else
       $F(i) \leftarrow n$ 
      i on stack for next pass
  endfor
  Send( topRow, rank-1 )
  Send( bottomRow, rank+1 )
  Recv( rowBelow, rank+1 )
  Recv( rowAbove, rank-1 )
Until F is not modified
    
```

Z denotes the original elevation.
 F denotes the pit filled elevation.
 n denotes lowest neighboring elevation
 i denotes the cell being evaluated

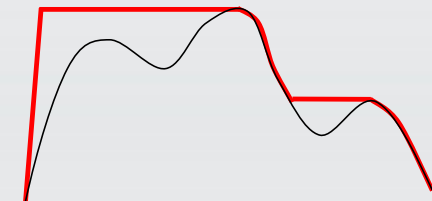
Iterate only over stack of changeable cells

Planchon, O. and F. Darboux, (2001),
[http://dx.doi.org/10.1016/S0341-8162\(01\)00164-3](http://dx.doi.org/10.1016/S0341-8162(01)00164-3).

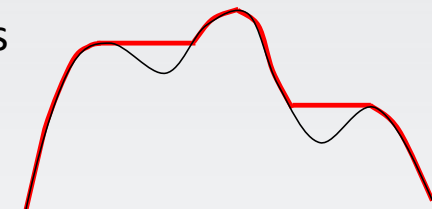
Initialization



1st Pass



2nd Pass



Continental-Scale HAND layer evaluated on ROGER supercomputer at NCSA using TauDEM



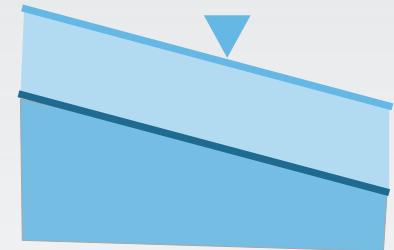
Catchments and Flowlines



Digital Elevation Model



**Height Above Nearest
Drainage (HAND)**
(relative elevation of land
surface cell above cell in
stream to which it flows)



Slide from David Maidment

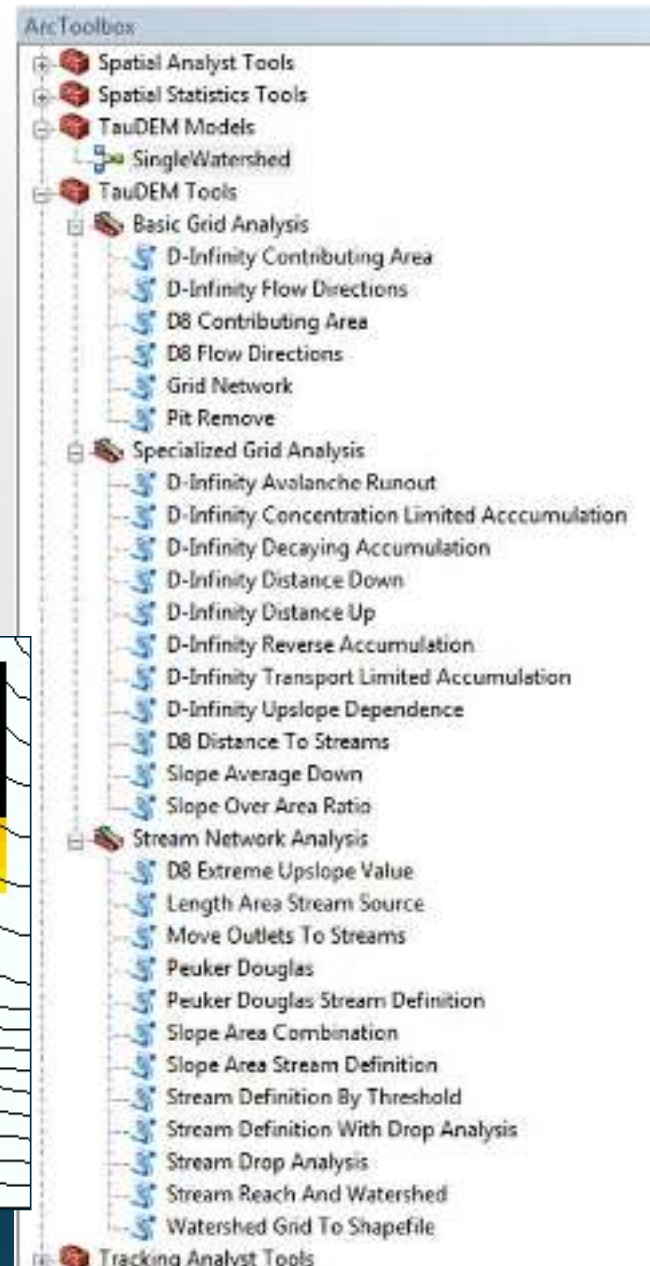
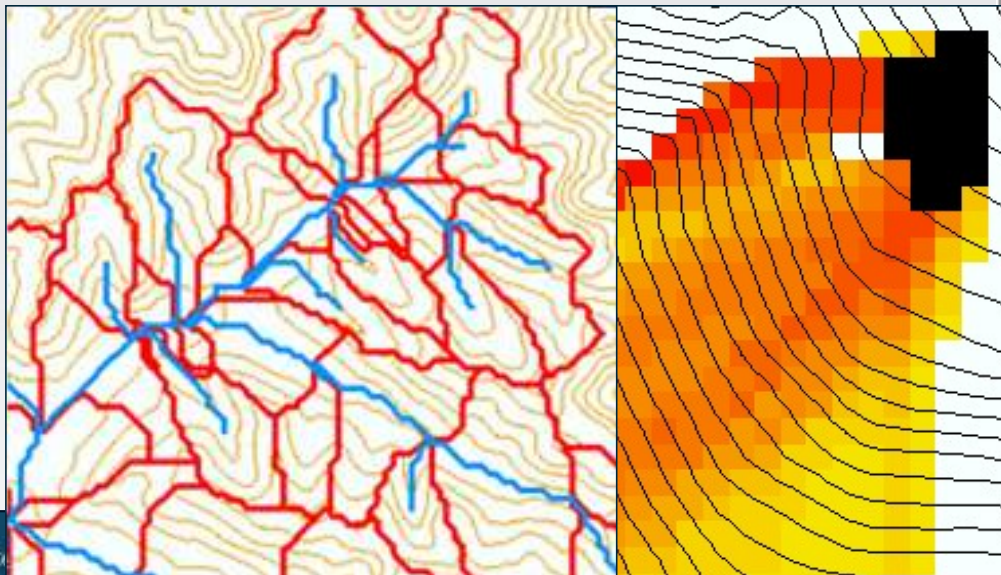


Utah Water Research Laboratory
Utah State University

<http://hydrology.usu.edu>

TauDEM

- D-infinity and D8 flow direction models
- Includes multiple hydrological proximity measures such as distance to stream which measured vertically gives height above the nearest drainage (HAND)
- Uses parallel algorithms for processing large datasets
- Includes Toolbox for ArcGIS
- Open source platform independent C++ command line executables for each function
- Deployed as an ArcGIS Toolbox with python scripts that drive command line executables



TauDEM Programming

- C++ Command Line Executables that use MPI
- Use GDAL/OGR library to read and write datasets in open standard formats for exchange with other programs
- ArcGIS Python Script Tools
- Python validation code to provide file name defaults
- Shared as ArcGIS Toolbox

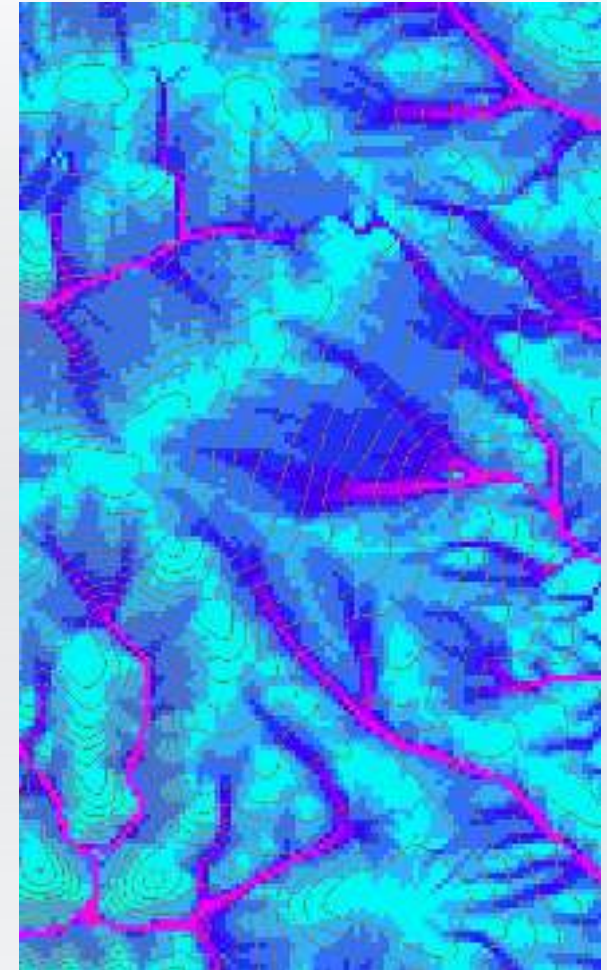
Any
computer

Requires
ArcGIS

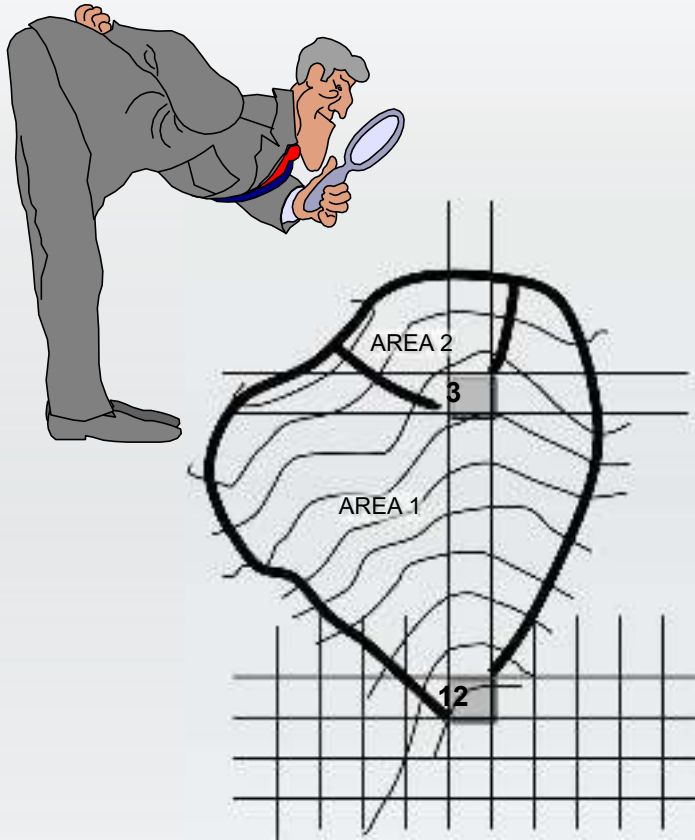


Conclusions

- Starting from a simple grid DEM, a rich set of data structures and information useful for hydrologic analysis can be derived
- Empirical geomorphology “laws” offer approaches objective selection of channel network delineation threshold
- Terrain surface derivatives enhanced by use of D_{∞} model.
- Flow algebra a general “recipe” for terrain flow related modeling.
- The height above nearest drainage approach provides a way to rapidly approximate real time flood inundation and approximate reach scale hydraulic properties
- Future challenges and opportunities remain as to how to leverage increasingly high resolution DEMs and emerging cloud and high-performance computing capabilities



Questions



Some References

- Tarboton, D. G., (1997), "A New Method for the Determination of Flow Directions and Contributing Areas in Grid Digital Elevation Models," *Water Resources Research*, 33(2): 309-319, <http://doi.org/10.1029/96WR03137>.
- Tarboton, D. G. and D. P. Ames, (2001), "Advances in the mapping of flow networks from digital elevation data," *World Water and Environmental Resources Congress*, Orlando, Florida, May 20-24, ASCE, [http://dx.doi.org/10.1061/40569\(2001\)166](http://dx.doi.org/10.1061/40569(2001)166).
- Tesfa, T. K., D. G. Tarboton, D. W. Watson, K. A. T. Schreuders, M. E. Baker and R. M. Wallace, (2011), "Extraction of hydrological proximity measures from DEMs using parallel processing," *Environmental Modelling & Software*, 26(12): 1696-1709, <http://dx.doi.org/10.1016/j.envsoft.2011.07.018>.
- Zheng, X., D. G. Tarboton, D. R. Maidment, Y. Y. Liu and P. Passalacqua, (2018), "River Channel Geometry and Rating Curve Estimation Using Height above the Nearest Drainage," *JAWRA Journal of the American Water Resources Association*, 54(4): 785-806, <http://doi.org/10.1111/1752-1688.12661>

Others at: <https://hydrology.usu.edu/dtarb/tarpubs.htm>

Thank You !

